Unravelling long-term behaviour in historic geophysical data sets

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Sodankylä Geophysical Observatory

- → First observations during theInternational Polar Year 1882/83.
- ⇒ Established 1913.
- Finland independent from Russia in 1917.
- → Part of University of Oulu since 1997.
- → Oldest scientific research institute in Northern Finland.



Where we are...



F-REGION THERMOSPHERE AURORA **E-REGION MESOSPHERE STRATOSPHERE** TROPOSPHERE **Electron Temperature Density**

Greenhouse Cooling

Doubling of [CO₂] and [CH₄]

cools

Mesosphere by 10 K and Thermosphere by 50 K.

MOSphere shrinks.

Layer of maximum electron density lowers by 15-20 km.

Greenhouse high up?

- Model results, assuming doubling of CO₂ and CH₄:
- Stratopause cools by 8 K, stratosphere by 15 K.
 (Brasseur & Hitchman, 1988)
- Mesosphere and thermosphere cool by 10 K and 50 K, respectively. (Roble & Dickinson, 1989)
- F2-layer peak (hmF2) lowers by 15-20 km. (Rishbeth, 1990)
- Riometer absorption decreases.
 (Serafimov & Serafimova, 1992)
- Stratopause cools by 14 K, mesosphere by 8 K, thermosphere by 50 K.
 (Akmaev & Fomichev, 1998)

lonosonde

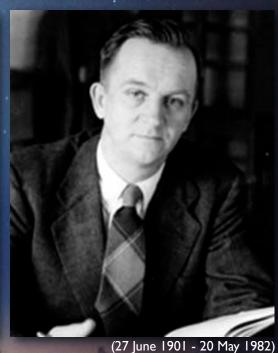
Gregory Breit

Merle Tuve

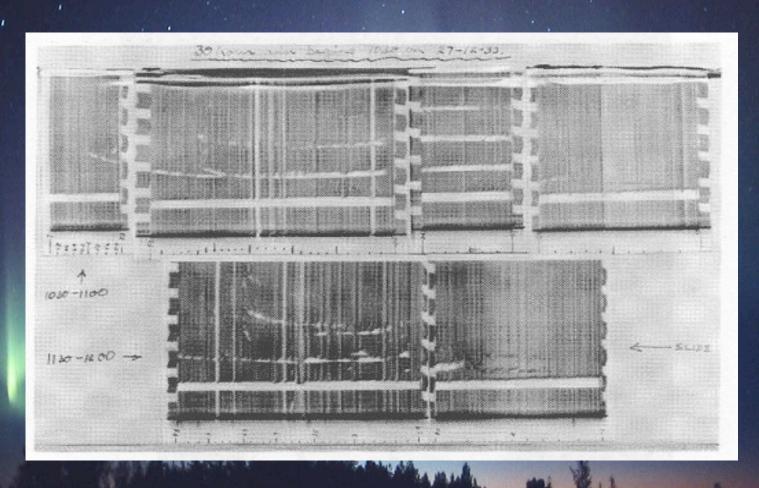
G Breit and MA Tuve, A radio method of estimating the height of the conducting layer, Nature, 116, p. 357, 1925.



(14 July 1899 - 11 September 198



Early Slough Ionogram



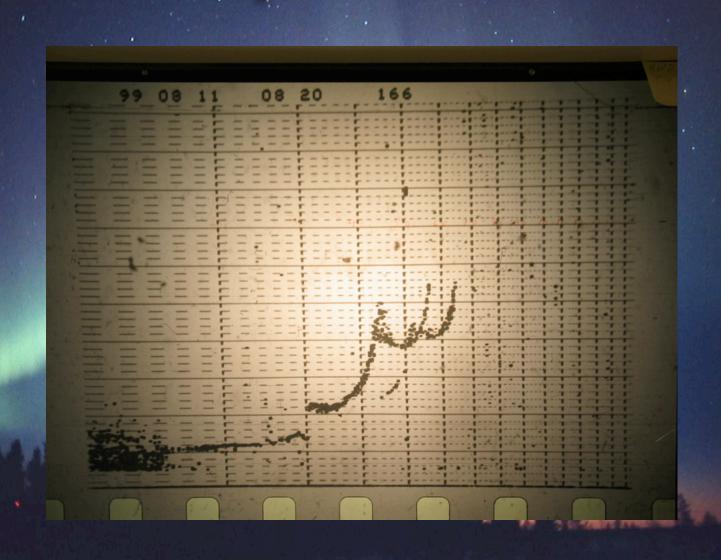
Radio Research Station Slough, Buckinghamshire 27th December 1933, 10:30-11:00 UTC and 11:30-12:00 UTC.

Sodankylä lonosonde

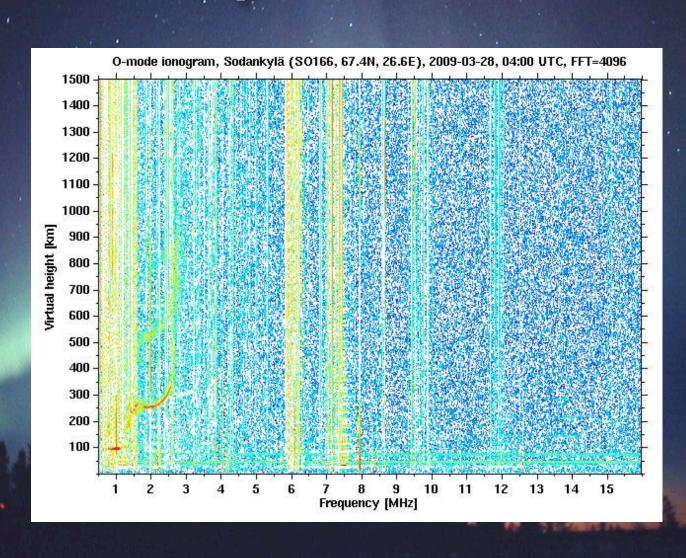
- Sodankylä ionosonde measurements began Ist August 1957.
- Until Nov 2005: I sounding per 30 min.
- Until Mar 2007: I sounding per 10 min.
- IPY (Apr '07-Mar '08): I sounding per minute.
- April 2008: we forgot to turn off IPY mode.
- Millionth ionogram: before May 2007.
- High data quality: first 800.000+ ionograms were analysed by the very same person!



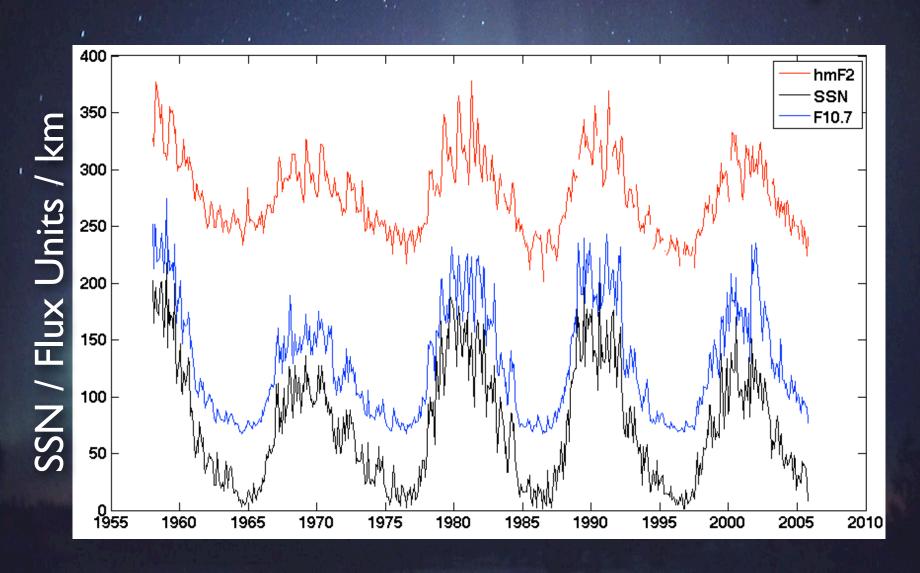
Sodankylä lonosonde



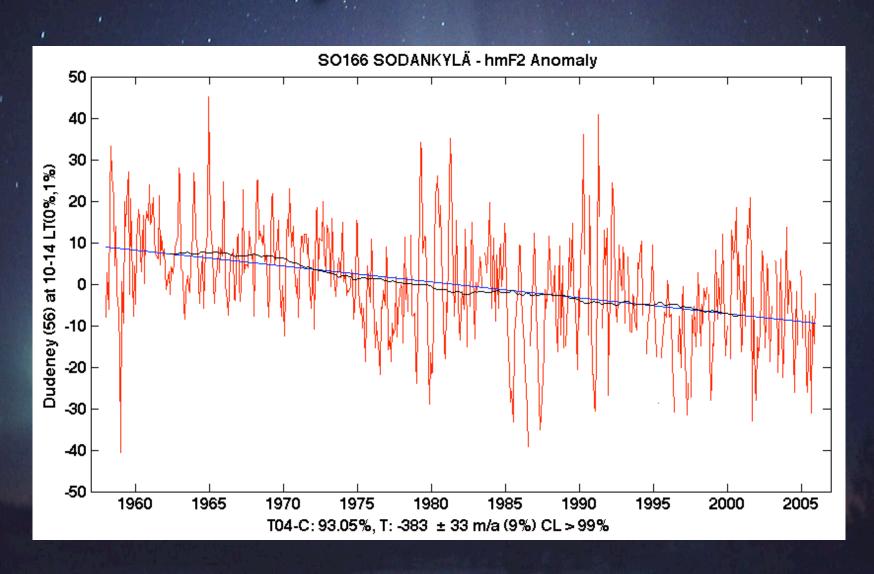
Sodankylä lonosonde



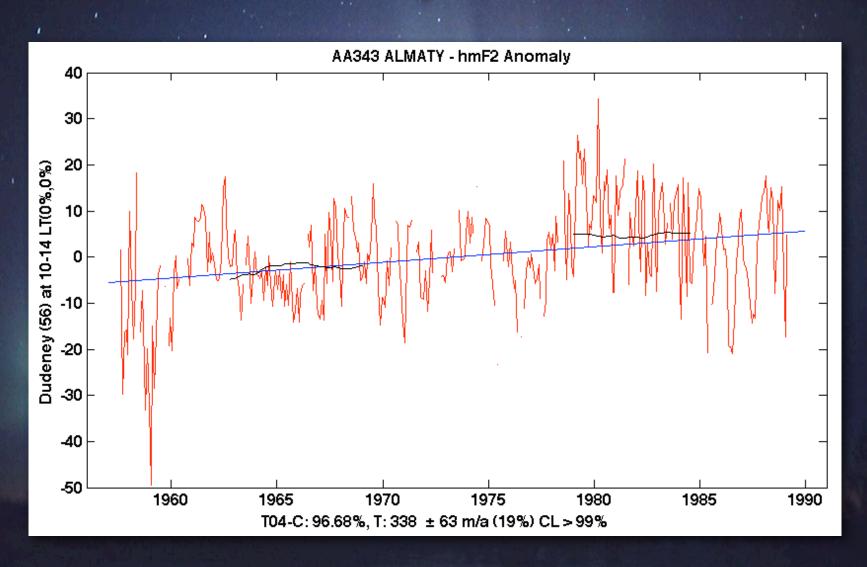
hmF2 & Solar Activity



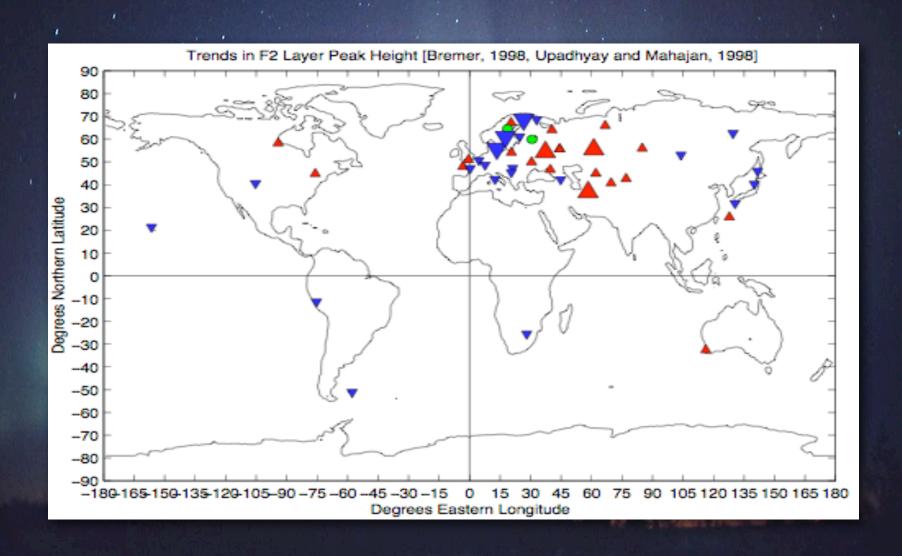
Sodankylä hmF2 Trend



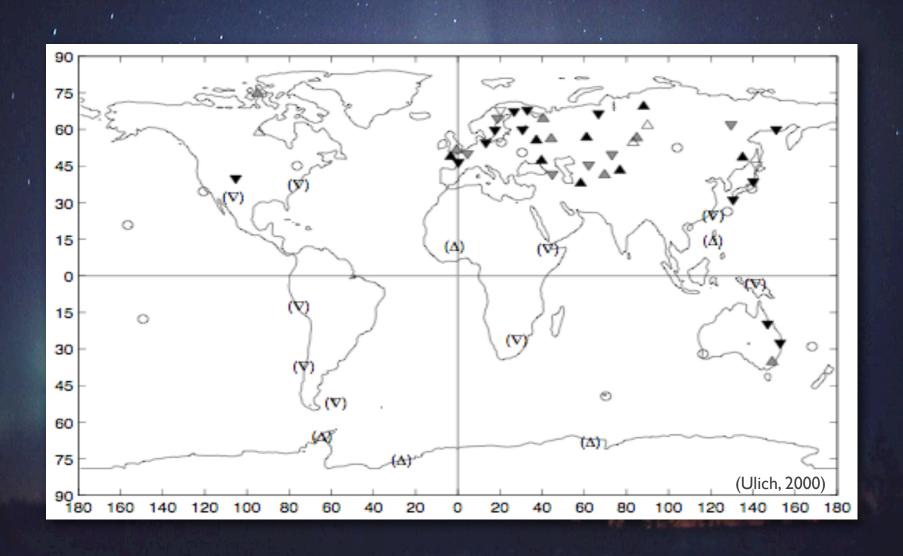
Almaty hmF2



hmF2 Trends



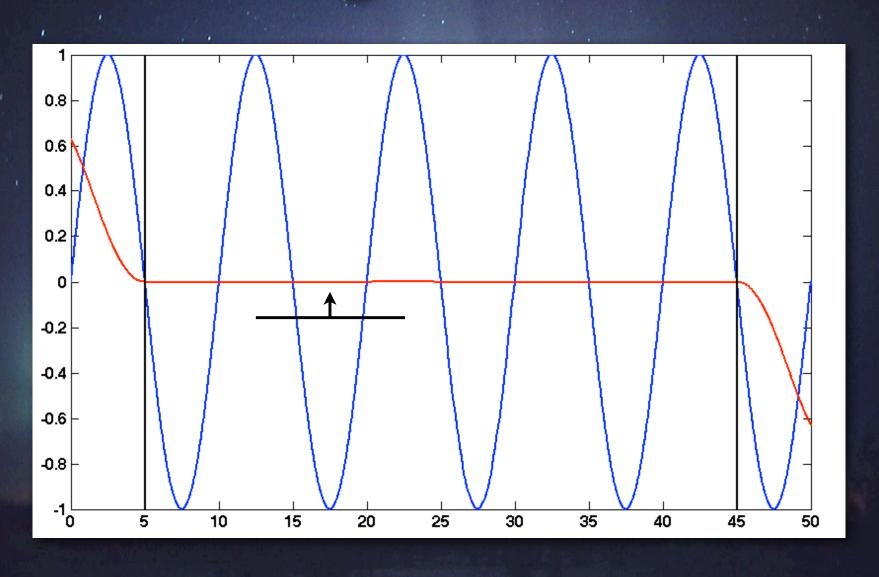
Global hmF2 Trends





- Data resolution (h, 3-h, day, month(?), ...)
- Low-pass filtering or polynomial fitting...

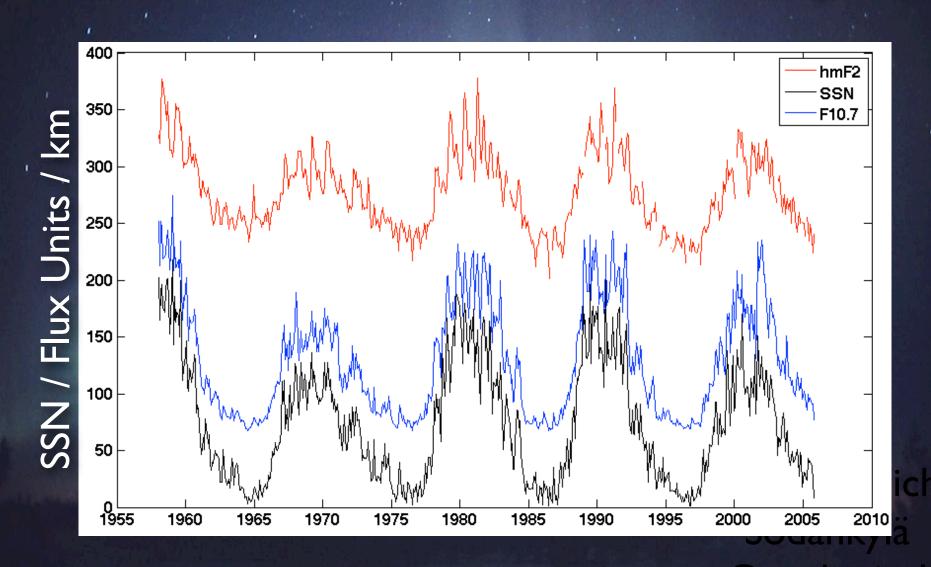
Running Mean Filter



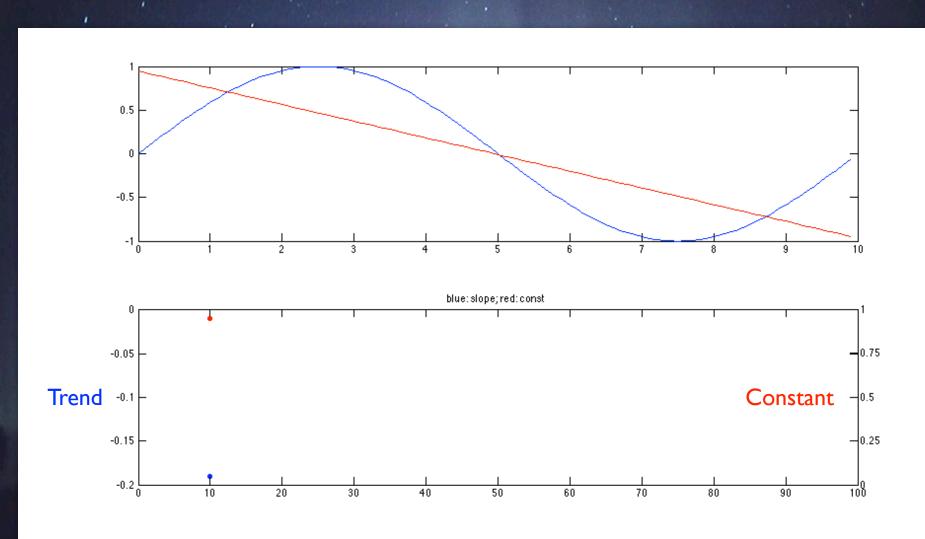
Problems

- Data resolution (h, 3-h, day, month(?), ...)
- Low-pass filtering or polynomial fitting...
- Removal of underlying (cyclic) variability:
 - Choice of proxy (sinusoid, SSN, Group SSN, F10.7 (adj./obs.), Ly-α, Mg II, E10.7, ...)
 - Resolution of proxy: compatibility with data

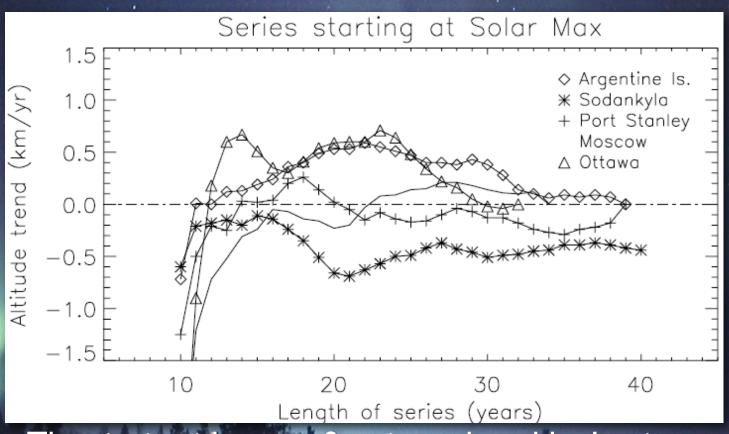
hmF2 & Solar Activity



Ringing



Ringing



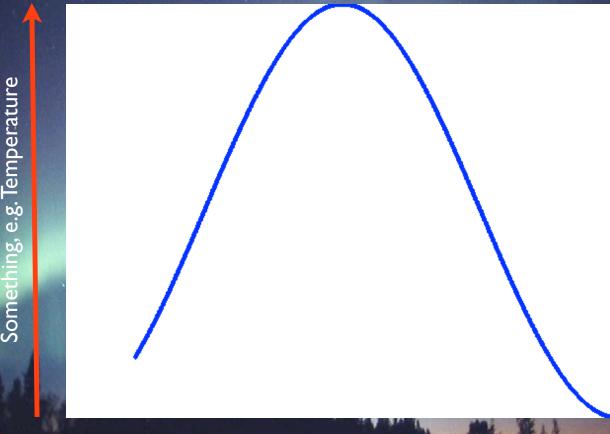
The ringing idea was first introduced by Jarvis et al., 2002. The plots shown here are from a follow-up paper by Clilverd et al., 2003.



- Data resolution (h, 3-h, day, month(?), ...)
- Low-pass filtering or polynomial fitting...
- Removal of underlying (cyclic) variability: ...
- Data gaps

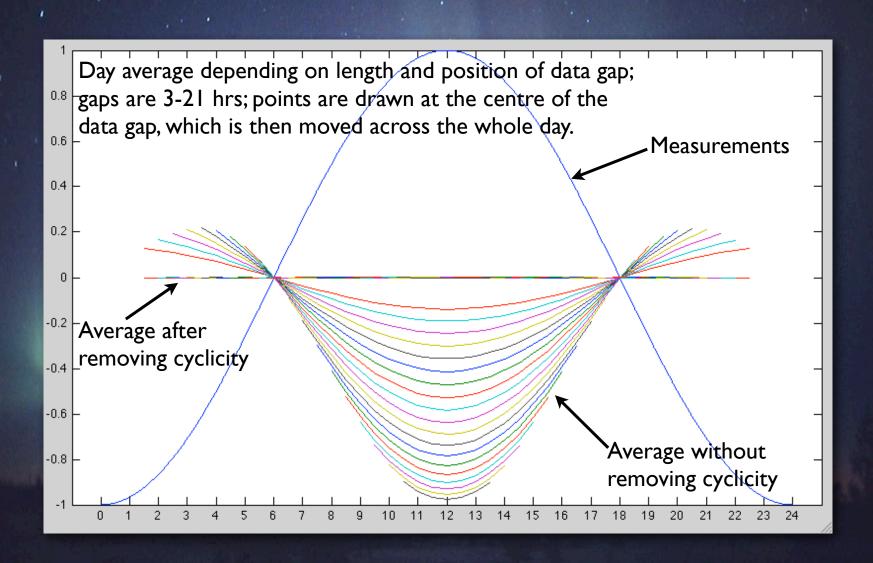
Example: Data Gaps

e.g. Temperature

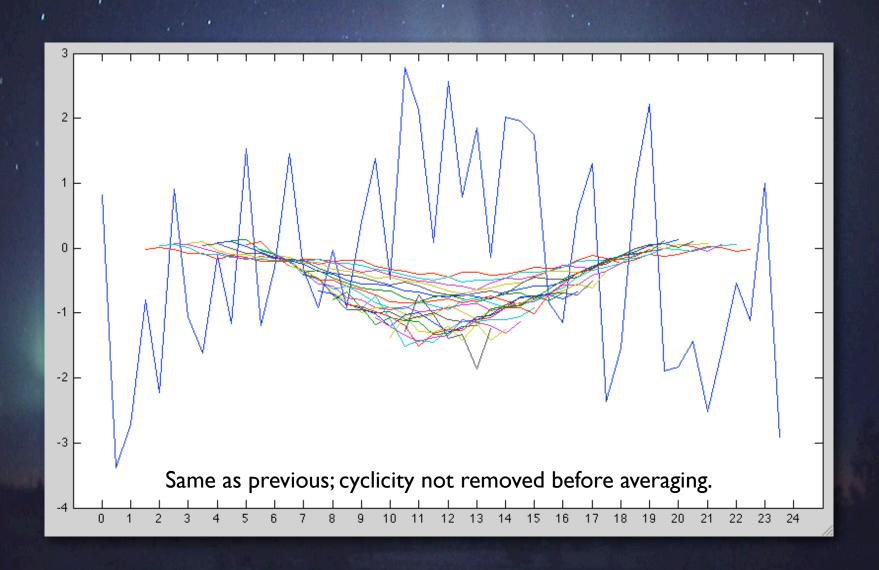


Time, e.g. I day, resolution 1/min

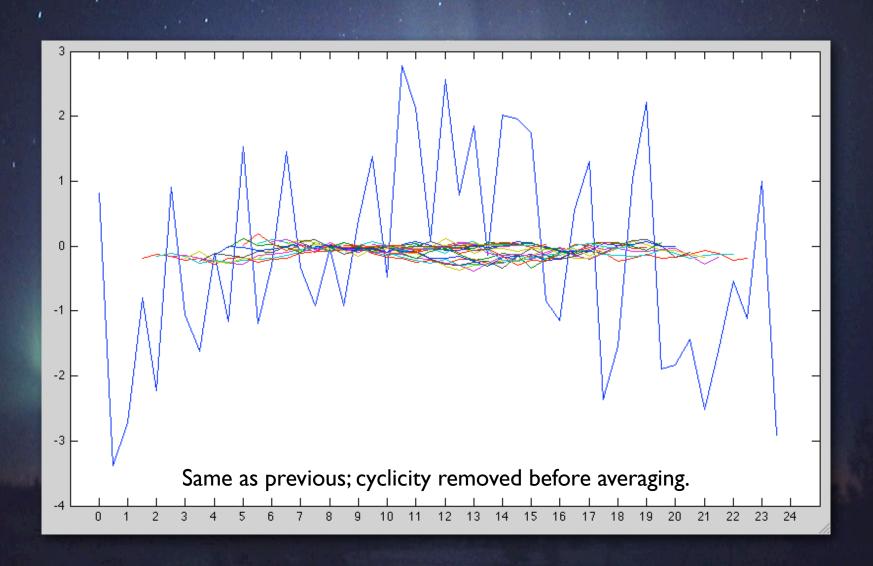
Data Gaps



Data Gaps



Data Gaps



Problems

- Data resolution (h, 3-h, day, month(?), ...)
- Low-pass filtering or polynomial fitting...
- Removal of underlying (cyclic) variability: ...
- Data gaps
- Measurement errors
- Mathematics of trend detection
 - stepwise or multi-parameter fit
 - error propagation

Making models

• Base functions of the model(s) are, e.g.:

$$\begin{array}{lll} m_i = \epsilon_i & -> \text{measurement errors} \\ + x_1 & -> \text{constant} \\ + x_2t_i & -> \text{sampling times} \\ + x_3F_{10.7}(t_i) & -> \text{solar activity} \\ + x_4Ap(t_i) & -> \text{geomagnetic activity} \\ + x_5\sin(2\pi t_i) & -> \text{annual variation} \\ + x_7\sin(4\pi t_i) & -> \text{semi-annual variation} \\ + x_8\cos(4\pi t_i) & -> \text{semi-annual variation} \\ + \end{array}$$

Modelling the data

The ionospheric property of interest is function of time and a number of other parameters. The model of the data is therefore

$$m(t) = \mathcal{F}(t, x_1, \dots, x_M)$$

where

$$\mathcal{F}(t, x_1, \dots, x_M) = \sum_{i=1}^M x_i f_i(t)$$

The actual measurements m_i observed at time t_i are equal to the model plus some measurement error ϵ_i

$$m_i = \mathcal{F}(t_i, x_1, \dots, x_M) + \varepsilon_i$$

Inverse problem I

This can be expressed as a matrix equation. Usually there are many more data points than unknowns x_i and the problem is over-determined:

$$\begin{pmatrix} m_1 \\ m_2 \\ \vdots \\ m_N \end{pmatrix} = \begin{pmatrix} f_1(t_1) & f_2(t_1) & \cdots & f_M(t_1) \\ f_1(t_2) & f_2(t_2) & \cdots & f_M(t_2) \\ \vdots & \vdots & \ddots & \vdots \\ f_1(t_N) & f_2(t_N) & \cdots & f_M(t_N) \end{pmatrix} \cdot \begin{pmatrix} x_1 \\ x_2 \\ \vdots \\ x_M \end{pmatrix} + \begin{pmatrix} \varepsilon_1 \\ \varepsilon_2 \\ \vdots \\ \varepsilon_N \end{pmatrix}$$

In other words:

$$\mathbf{m} = \mathbf{A} \cdot \mathbf{x} + \boldsymbol{\varepsilon}$$

Inverse problem II

Measurements and theory are weighted by the measurement errors:

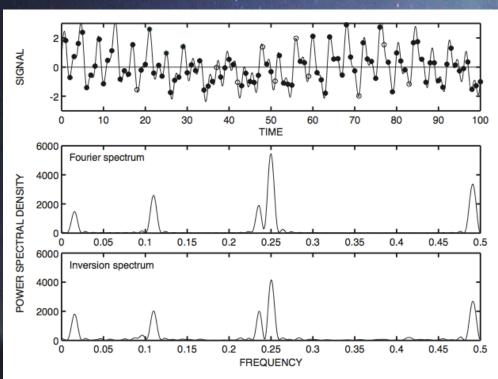
$$B_{ij} := \frac{A_{ij}}{\varepsilon_i}$$
 and $b_i := \frac{m_i}{\varepsilon_i}$

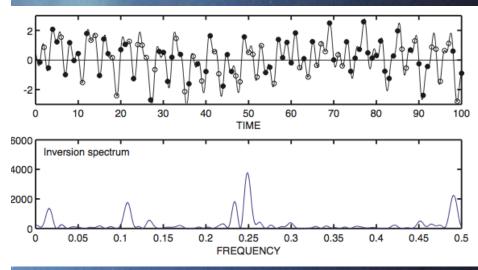
The solution is the vector **x**, which minimises the following expression:

$$\chi^2 = |\mathbf{B} \cdot \mathbf{x} - \mathbf{b}|^2$$

We are left with a general least squares problem. Solving this results in the most probable solution for **x**.

Signal Spectrum by Stochastic Inversion

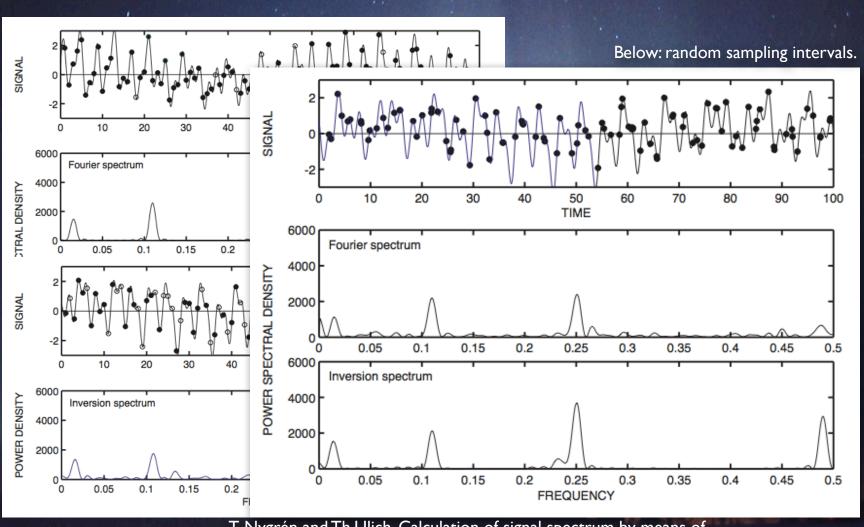




Left: 100 pts for Fourier, 90 for inversion. Above: 59 pts.

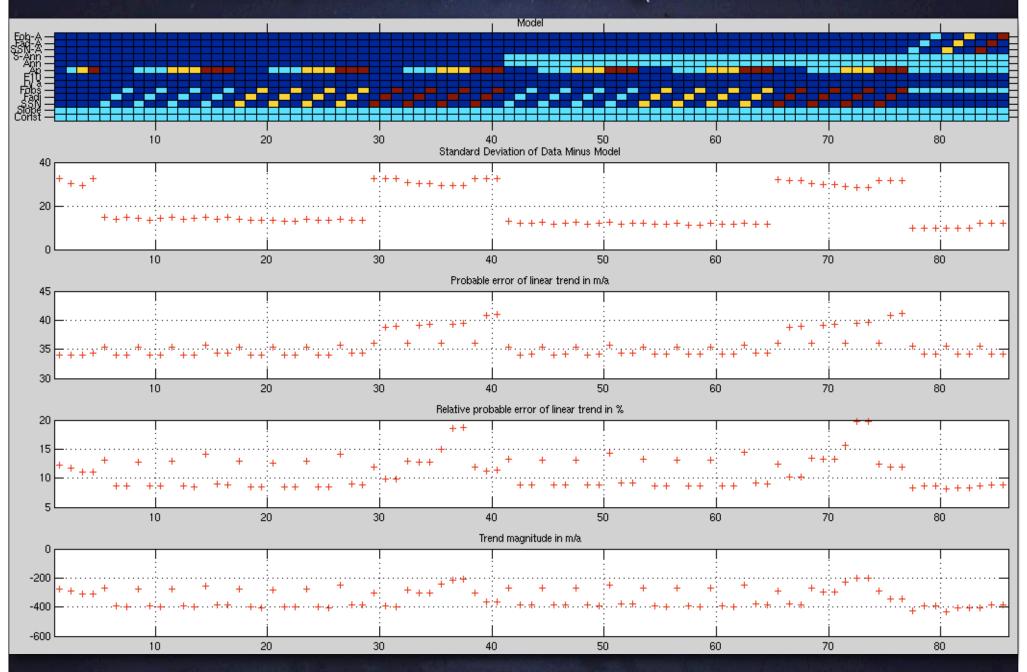
T. Nygrén and Th. Ulich, Calculation of signal spectrum by means of stochastic inversion, Ann. Geophys., 28, 1409-1418, 2010.

Signal Spectrum by Stochastic Inversion



T. Nygrén and Th. Ulich, Calculation of signal spectrum by means of stochastic inversion, Ann. Geophys., 28, 1409-1418, 2010.

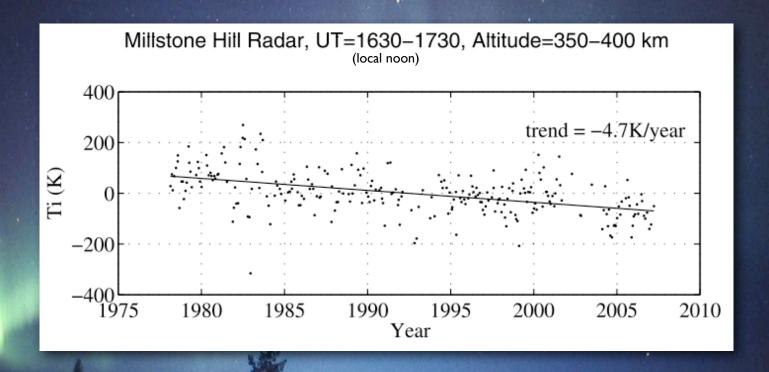
Sodankylä F2-layer peak height hmF2



Trends in other Observations

Height	Method	Parameter	Trend	Reference
in km			per Year	
75	Sounding rocket	Temperature	-0.6 K	Kokin and Lysenko, 1994
70	Sounding rocket	Temperature	-0.7 K	Golitsyn et al., 1996
60-70	Lidar	Temperature	-0.4 K	Hauchecorne et al., 1991
60	Sounding rocket	Temperature	-0.4 K	Golitsyn et al., 1996
60	Sounding rocket	Temperature	-0.33 K	Keckhut et al., 1999
50-60	Lidar	Temperature	-0.25 K	Aikin et al., 1991
50	Sounding rocket	Temperature	-0.25 K	Golitsyn et al., 1996
40	Sounding rocket	Temperature	-0.1 K	Golitsyn et al., 1996
30-60	Sounding rocket	Temperature	-0.17 K	Dunkerton et al., 1998
30-50	Sounding rocket	Temperature	-0.17 K	Keckhut et al., 1999
30	Sounding rocket	Temperature	-0.1 K	Golitsyn et al., 1996
25	Sounding rocket	Temperature	-0.1 K	Golitsyn et al., 1996
25	Sounding rocket	Temperature	-0.11 K	Keckhut et al., 1999

Direct F-Region Temperature



Long-term temperature trends in the ionosphere above Millstone Hill

J. M. Holt1 and S. R. Zhang1

GEOPHYSICAL RESEARCH LETTERS, VOL. 35, L05813, doi:10.1029/2007GL031148, 2008

Conclusion

(the last one, I promise!)

- Definitely, there's long-term change in the ionosphere and thermosphere!
- The enhanced greenhouse effect is probably a part of it.
- Other (unknown?) processes are involved.
- Solution in modelling?
- We don't understand what's going on.
- Student exercise: Find out!

Conclusion

(I lied to you!)

lonsondes, originally deployed for monitoring ionospheric conditions for HF radio communication and for studying short-term events, are becoming useful in an environmental context.

They provide long-term measurements of our environment!

Do not discontinue atmospheric observations at a time of climate change!