

An Introduction to the Terrestrial Ionosphere

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1 Local Photochemistry and Energetics

- Photochemistry
- Energetics

2 Transport

- Thermal Plasma Transport
- Suprathermal Particle Transport

3 Electrodynamics

- Dynamo Theory
- High Latitude Electrodynamics

Hydrostatic Equilibrium

Vertical structure of neutral atmosphere determined by balance between **pressure gradient** and **gravity**.

$$\nabla p = -mng$$

$$\frac{d}{dz}nk_B T = -mng$$

In the simplest case of constant T Define the scale height:

$$k_B T \frac{d}{dz}n = -mng$$

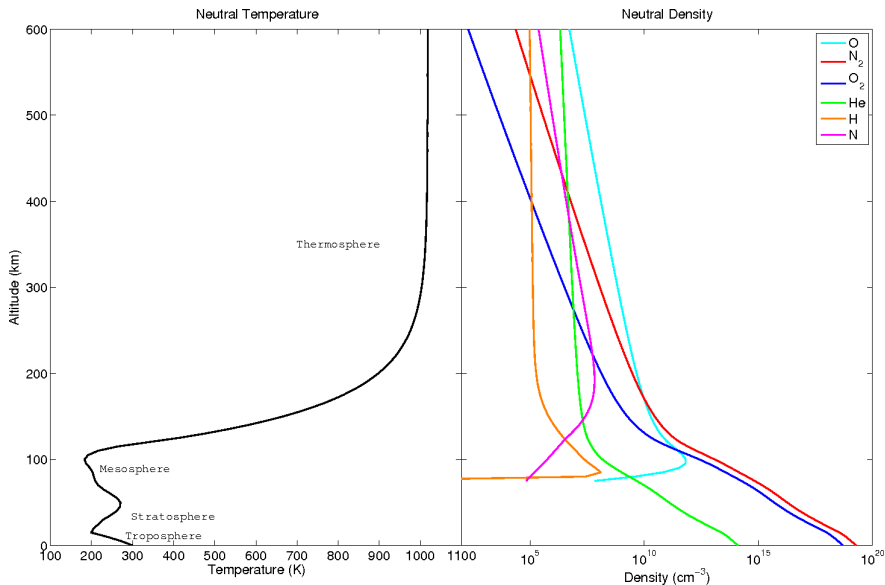
$$H = \frac{k_B T}{mg}$$

$$\frac{1}{n} \frac{dn}{dz} = -\frac{mg}{k_B T} = -\frac{1}{H}$$

Solution:

$$n(z) = n(z_0) \exp \left[-\frac{z - z_0}{H} \right]$$

The Neutral Atmosphere (according to NRLMSISE-00)



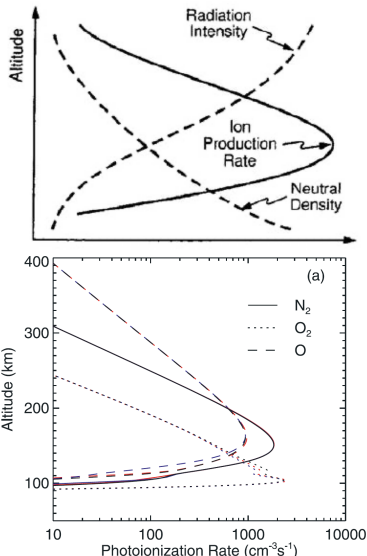
Absorption of EUV Radiation

Photoionization Rate:

$$P_i(s) = \int d\lambda \sum_n \sigma_n^{ion}(\lambda) N_n(s) I(s, \lambda)$$

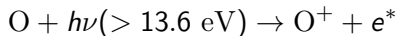
Attenuation of EUV Flux
(Lambert's Law):

$$\frac{dI}{ds} = - \sum_n \sigma_n^{abs}(\lambda) N_n(s) I(s, \lambda)$$

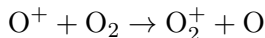
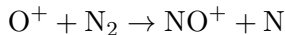


Basic Ionospheric Photochemistry

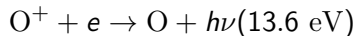
Photoionization:



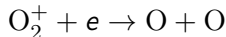
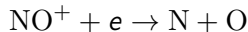
Atom-Ion Interchange:



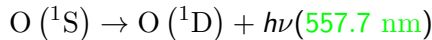
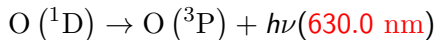
Radiative Recombination (SLOW):



Dissociative Recombination:



Airglow:



The Ionosphere and Thermosphere

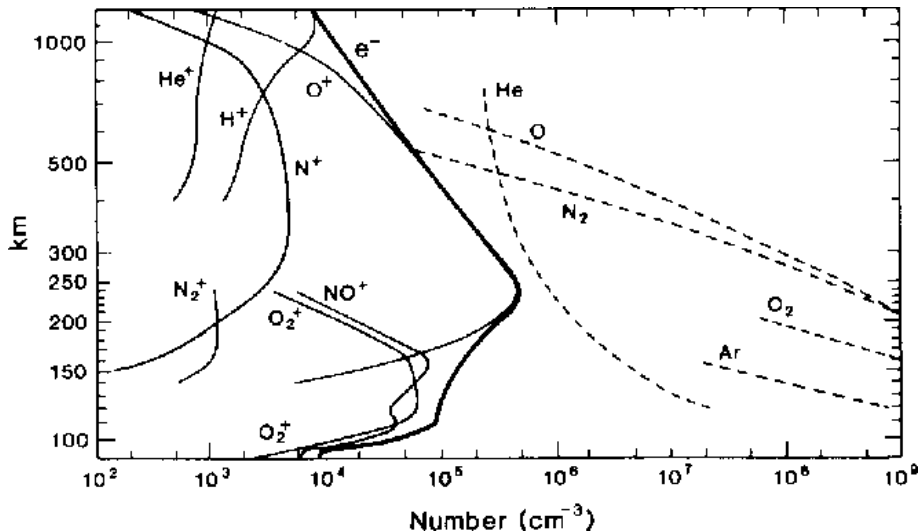
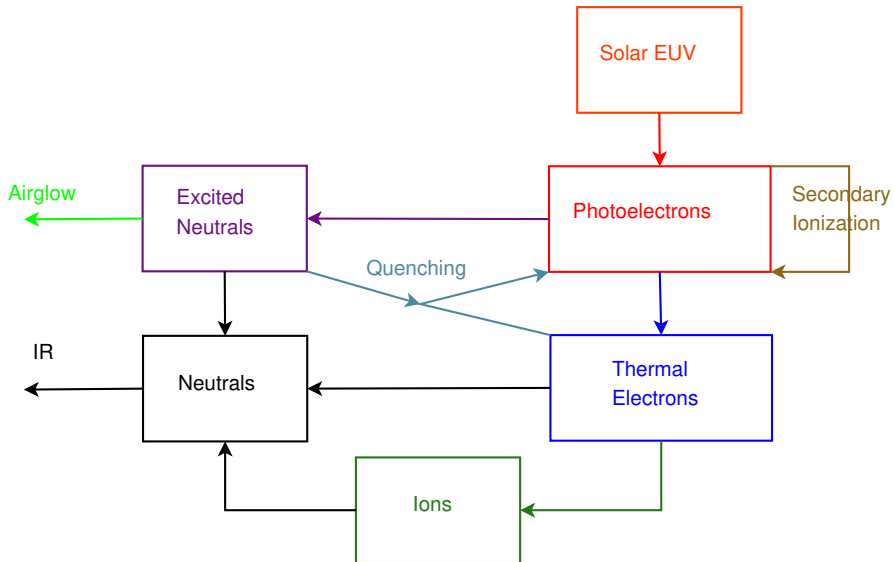
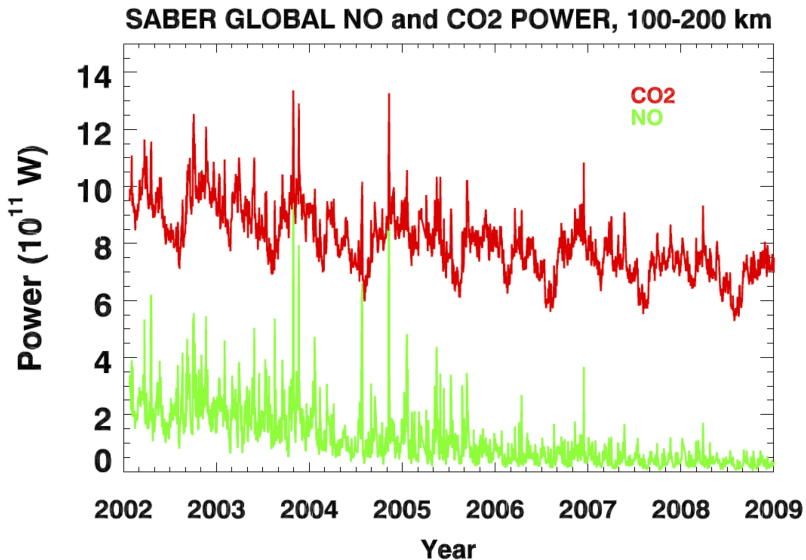


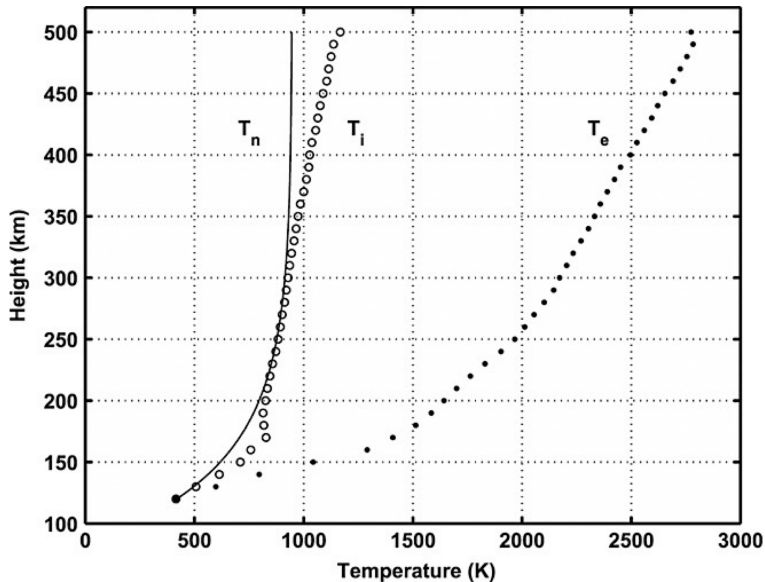
Diagram of Ionospheric Energetics



IR Cooling of Thermosphere

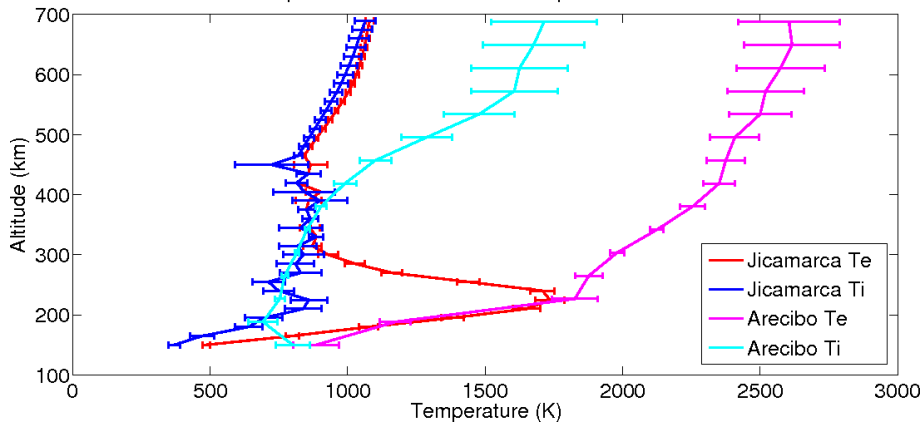


Typical Daytime Mid-latitude Temperature Profiles



Temperature Profiles at Different latitudes

Equatorial vs. Mid-latitude Temperature Profiles



Importance of the Magnetic Field

Dynamic Pressure: $mn u^2$

Thermal Pressure: $nk_B T = mn \left(\sqrt{\frac{k_B T}{m}} \right)^2$

Magnetic Pressure: $\frac{B^2}{2\mu_0} = mn \frac{1}{2} \left(\frac{B}{\sqrt{\mu_0 mn}} \right)^2$

Typical Numbers for the ionosphere:

$u \approx 100$ m/s Midlatitude Ionosphere

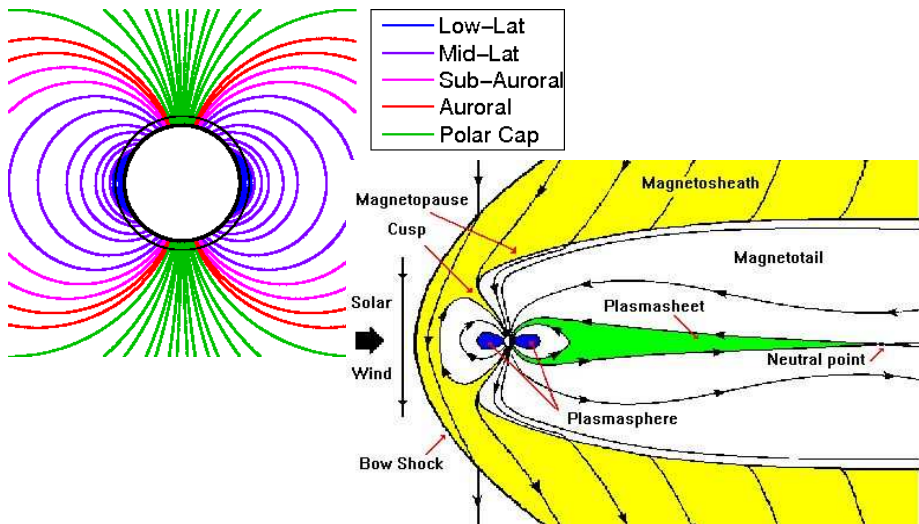
$u \approx 20$ km/s H^+ Outflow in Polar Wind

$T \approx 2000$ K $\rightarrow \sqrt{\frac{k_B T}{m}} = \begin{cases} 1 \text{ km/s} & \text{for } O^+ \\ 4 \text{ km/s} & \text{for } H^+ \end{cases}$

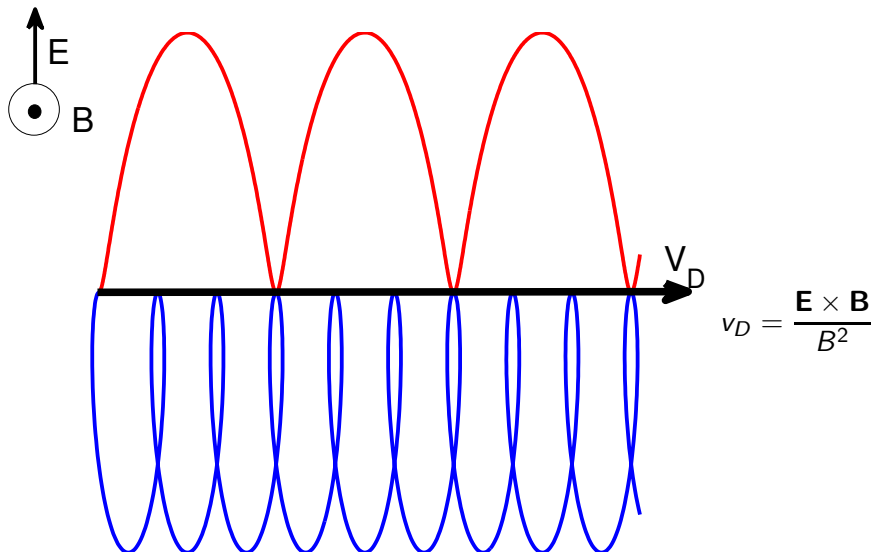
$B \approx 3 \times 10^{-5}$ T, $n \approx 10^{11} \text{ m}^{-3} \rightarrow \frac{B}{\sqrt{\mu_0 mn}} \approx 500 \text{ km/s}$

Transport parallel and perpendicular to B are fundamentally different.

Magnetic Structure of the Ionosphere and Magnetosphere



Perpendicular Transport: The $\mathbf{E} \times \mathbf{B}$ Drift



Moments of the Boltzmann Equation

Boltzmann Equation:

$$\frac{\partial f}{\partial t} + \mathbf{v} \cdot \frac{\partial f}{\partial \mathbf{x}} + \left[\frac{e}{m} (\mathbf{E} + \mathbf{v} \times \mathbf{B}) + \mathbf{g} \right] \cdot \frac{\partial f}{\partial \mathbf{v}} = \left. \frac{\delta f}{\delta t} \right|_{\text{collisions}}$$

Continuity Equation: Apply $\int d\mathbf{v}$ to Boltzmann Equation

$$\frac{\partial n}{\partial t} + \frac{\partial}{\partial \mathbf{x}} \cdot (n\mathbf{u}) = \frac{\delta n}{\delta t}$$

Momentum Equation: Apply $\int m\mathbf{v}d\mathbf{v}$ to Boltzmann Equation

$$\frac{\partial}{\partial t} (m n \mathbf{u}) + \frac{\partial}{\partial \mathbf{x}} \cdot (m n \mathbf{u} \mathbf{u} + \mathbf{P}) = n e (\mathbf{E} + \mathbf{u} \times \mathbf{B}) + n m \mathbf{g} + \frac{\delta M}{\delta t}$$

Energy Equation: Apply $\int \frac{1}{2} m v^2 d\mathbf{v}$ to Boltzmann Equation

$$\frac{\partial \epsilon}{\partial t} + \frac{\partial}{\partial \mathbf{x}} \cdot [\mathbf{u} \cdot (\epsilon \mathbf{I} + \mathbf{P}) + \mathbf{q}] = n e \mathbf{u} \cdot \mathbf{E} + n m \mathbf{u} \cdot \mathbf{g} + \frac{\delta E}{\delta t}$$

Closing the System of Transport Equations

- Easiest way: Assume an isotropic Maxwellian distribution $\rightarrow$ 5-moment approximation

$$\mathbf{P} \rightarrow p\mathbf{I} \quad \mathbf{q} \rightarrow 0$$

Energy equation reduces to the adiabatic gas law $\frac{D}{Dt} \left(\frac{p}{n^\gamma} \right) = 0$

- Hard way: Assume more complicated distributions (e.g. Maxwellians times truncated series expansions). The 8-, 10-, 13-, 16-, and 20-moment equations are derived this way.
- Middle Ground: Assume higher moments are small and derive steady state limits of high-moment transport equations

$$\mathbf{q} = -\kappa \cdot \nabla T$$

Thermal Conduction

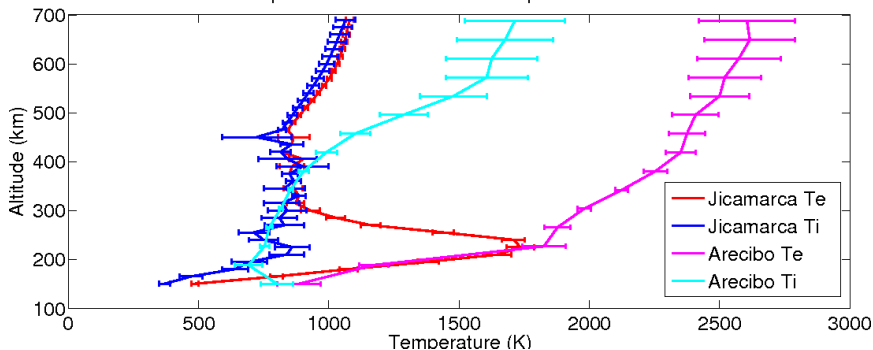
In the F -region $\kappa = \kappa_{\parallel} \hat{\mathbf{b}}\hat{\mathbf{b}} \rightarrow \mathbf{q} = -\kappa_{\parallel} \nabla_{\parallel} T \hat{\mathbf{b}}$

For a fully ionized plasma $\kappa_{\parallel} = 7.7 \times 10^5 T_e^{5/2} \text{ eV cm}^{-2} \text{ s}^{-1} \text{ K}^{-1}$

Parallel Temperature Equation:

$$\frac{\partial T}{\partial t} + u_{\parallel} \nabla_{\parallel} T + \frac{2}{3} T \nabla_{\parallel} \cdot \mathbf{u} - \frac{2}{3} \frac{1}{n k_B} \nabla_{\parallel} \cdot \kappa_{\parallel} \nabla_{\parallel} T = \frac{2}{3} \frac{1}{n k_B} (Q - L)$$

Equatorial vs. Mid-latitude Temperature Profiles



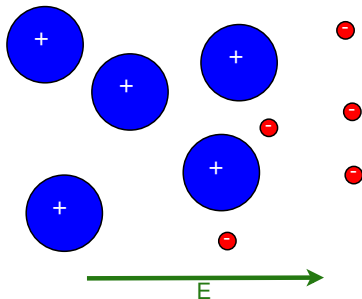
Ambipolar Electric Fields and Ambipolar Diffusion

Steady state parallel electron momentum equation:

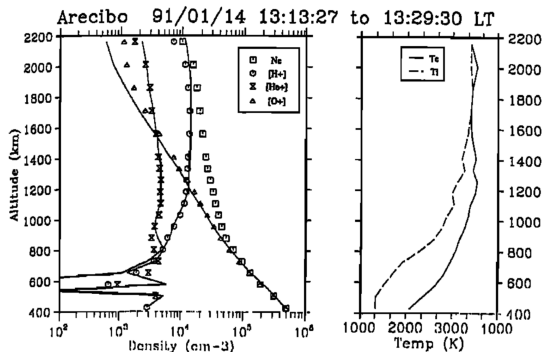
$$m_e \left[\cancel{\frac{\partial}{\partial t} (n_e u_e)} + \cancel{\nabla_{\parallel} \cdot (n_e u_e^2)} \right] = -\nabla_{\parallel} p_e - n_e e E_{\parallel} \longrightarrow E_{\parallel} = -\frac{1}{en_e} \nabla_{\parallel} p_e$$

Substitute into parallel ion momentum equation:

$$m_i \left[\frac{\partial}{\partial t} (n_i u_i) + \nabla_{\parallel} \cdot (n_i u_i^2) \right] = -\nabla_{\parallel} p_i - \frac{n_i}{n_e} \nabla_{\parallel} p_e - m_i n_i g_{\parallel} - m_i n_i \sum_j \nu_{ij} (u_i - u_j)$$



Topside Structure



Major Ion Scale Height

$$H_p = \frac{k_B (T_e + T_i)}{m_i g}$$

Minor Ion Scale Height

$$H_j = \left(\frac{m_j g}{k_B T_j} - \frac{q_j}{e} \frac{T_e}{T_j} \frac{1}{H_p} \right)^{-1}$$

Erickson and Swartz (1994)

Energetic Electron Transport

Populations of electrons in the ionosphere:

- Thermal: $k_B T_e \sim 0.2$ eV
- Photoelectrons: mostly < 60 eV, peak energy flux at ~ 20 eV
- Soft Precipitation (e.g. cusp, polar rain): $100 - 1000$ eV
- Auroral Precipitation: > 1 keV

Simplified kinetic equations derived by

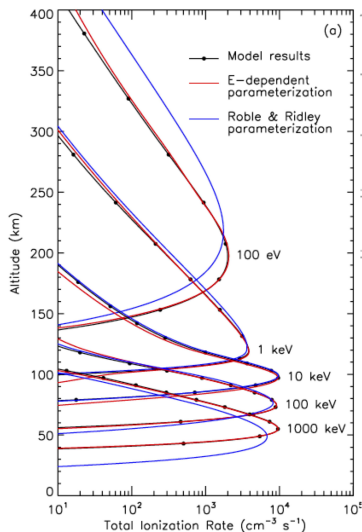
- Assuming suprathemal density $\ll$ thermal density
- Ignoring perpendicular transport
- Assuming gyrotropy (azimuthal symmetry about $\mathbf{B}$)
- Assuming steady state ($m_e \rightarrow 0$)

Simplest possible form is derived by additionally neglecting $E_{\parallel}$, $\frac{\partial B}{\partial s}$, and Coulomb collisions and assuming isotropic elastic collisions.

$$\mu \frac{\partial \Phi}{\partial s} = q + \sum_n \left\{ -[\sigma_{an}(\mathcal{E}) + \sigma_{en}(\mathcal{E})] N_n \Phi + \frac{\sigma_{en}(\mathcal{E})}{2} N_n \int_{-1}^1 \Phi(s, \mathcal{E}, \mu') d\mu' \right\}$$

This has the same mathematical form as a radiative transfer equation

Auroral Particle Deposition



Higher energy particles penetrate deeper into atmosphere.

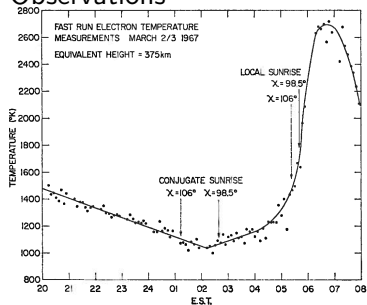


Fang et al. (2008)

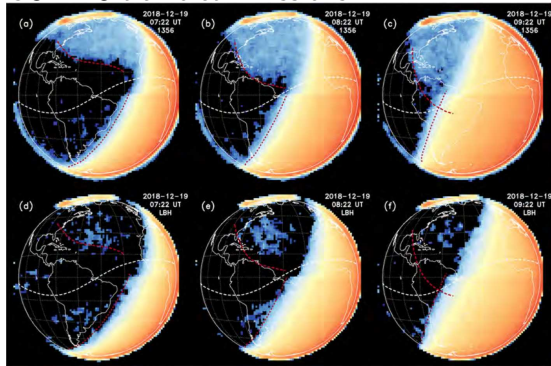
Conjugate Photoelectron Transport

Millstone Hill T_e

Observations



GOLD Ultraviolet Emissions



Solomon et al. 2020

Fundamentals of Ionospheric Electrodynamics

Electrostatic Limit of Maxwell's Equations:

$$\begin{aligned}\nabla \times \mathbf{B} &= \mu_0 \mathbf{J} + \frac{1}{c^2} \frac{\partial \mathbf{E}}{\partial t} && \longrightarrow \nabla \cdot \mathbf{J} = 0 \\ \nabla \times \mathbf{E} &= -\frac{\partial \mathbf{B}}{\partial t} && \longrightarrow \mathbf{E} = -\nabla \Phi\end{aligned}$$

Ohm's Law for the ionosphere:

$$\mathbf{J} = \sigma \cdot \mathbf{E} + \mathbf{J}_0$$

Putting everything together yields a boundary value problem:

$$\nabla \cdot \sigma \cdot \nabla \Phi = \nabla \cdot \mathbf{J}_0$$

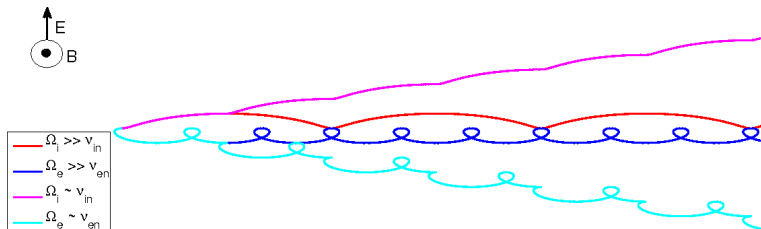
Ohm's Law for the Ionosphere

Steady-state momentum equation for each species (zero neutral wind case):

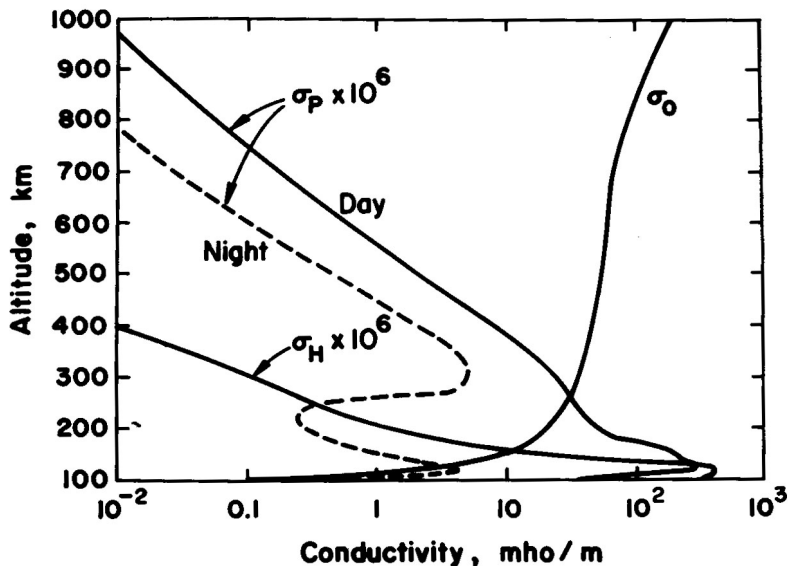
$$0 = n_{\alpha} q_{\alpha} (\mathbf{E} + \mathbf{u}_{\alpha} \times \mathbf{B}) - \nu_{\alpha n} m_{\alpha} n_{\alpha} \mathbf{u}_{\alpha}$$

Resulting Ohm's Law:

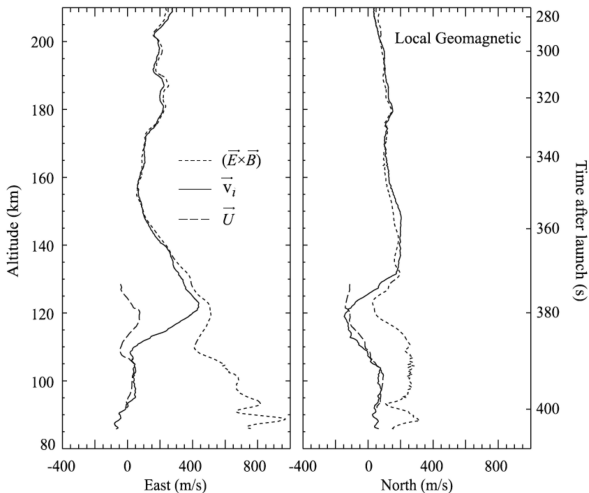
$$\mathbf{J} = \sum_{\alpha} n_{\alpha} q_{\alpha} \mathbf{u}_{\alpha} \longrightarrow \mathbf{J} = \begin{pmatrix} \sigma_P & -\sigma_H & 0 \\ \sigma_H & \sigma_P & 0 \\ 0 & 0 & \sigma_0 \end{pmatrix} \cdot \mathbf{E}$$



Conductivity Profiles



Ion Velocity Rotation

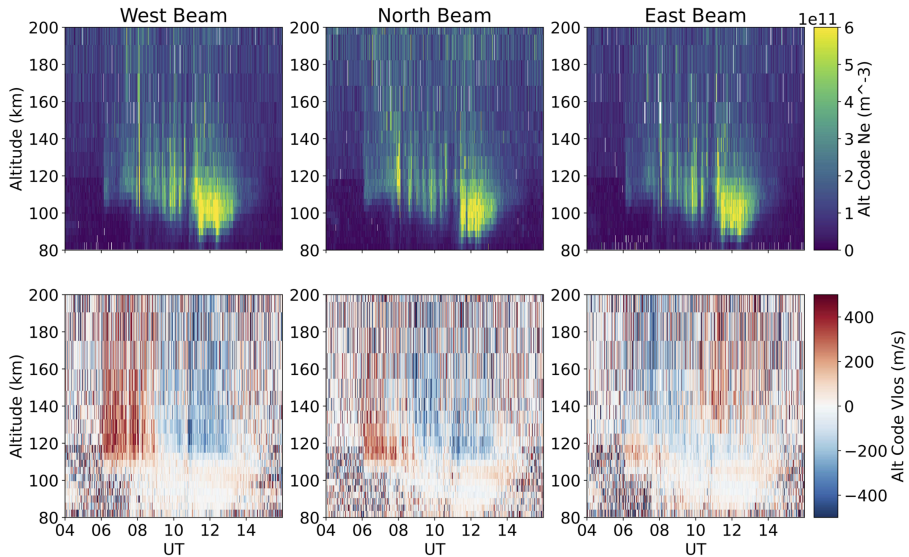


Rocket measurements
comparing

- $\vec{E}$ from double-probe
- $\vec{v}_i$ from ion imager
- $\vec{u}_n$ from TMA
chemical release

Sangalli et al. (2009)
doi:10.1029/2008JA013757

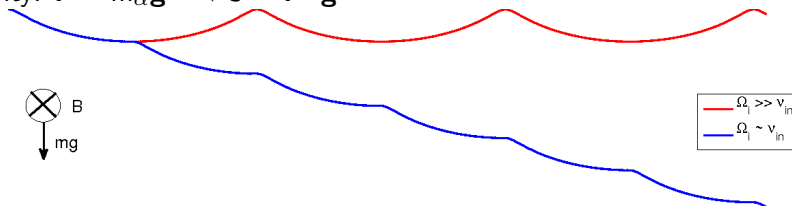
PFISR E-region Velocity Measurements



Other Kinds of Current

Substitute $\mathbf{F}$ for $q_\alpha \mathbf{E}$ in steady state momentum equation.

- Wind drag: $\mathbf{F} = \nu_{\alpha n} m_\alpha \mathbf{u}_n \longrightarrow \mathbf{J} = \sigma \cdot (\mathbf{u}_n \times \mathbf{B})$
- Gravity: $\mathbf{F} = m_\alpha \mathbf{g} \longrightarrow \mathbf{J} = \mathbf{\Gamma} \cdot \mathbf{g}$



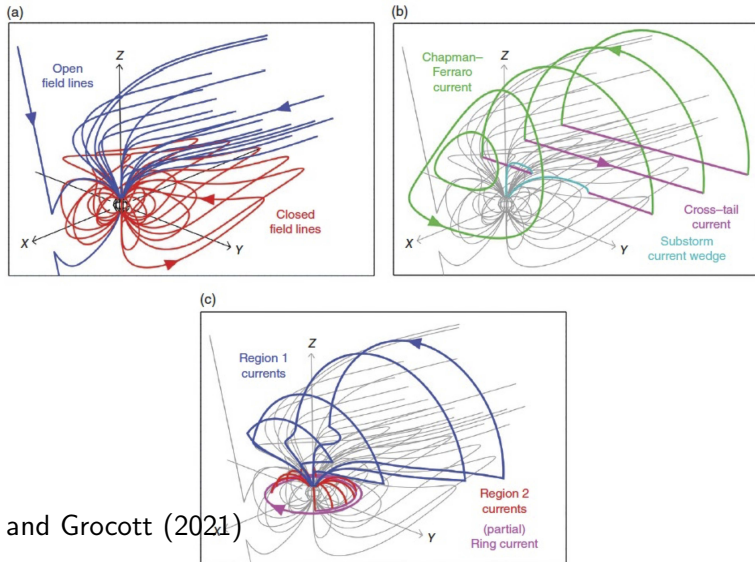
- Pressure Gradients (Diamagnetic Currents):

$$\mathbf{F} = -\frac{1}{n_\alpha} \nabla p_\alpha \longrightarrow \mathbf{J} = \mathbf{D} \cdot \nabla \sum_\alpha p_\alpha$$

Complete Dynamo Equation:

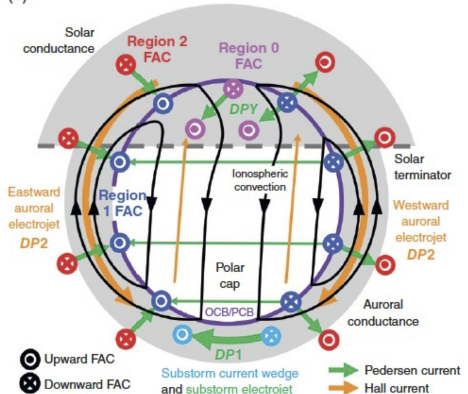
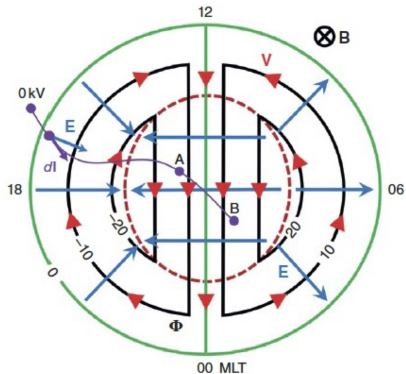
$$\nabla \cdot \sigma \cdot \nabla \Phi = \nabla \cdot \left(\sigma \cdot (\mathbf{u}_n \times \mathbf{B}) + \mathbf{\Gamma} \cdot \mathbf{g} + \mathbf{D} \cdot \nabla \sum_\alpha p_\alpha \right)$$

Current Systems in the Ionosphere and Magnetosphere

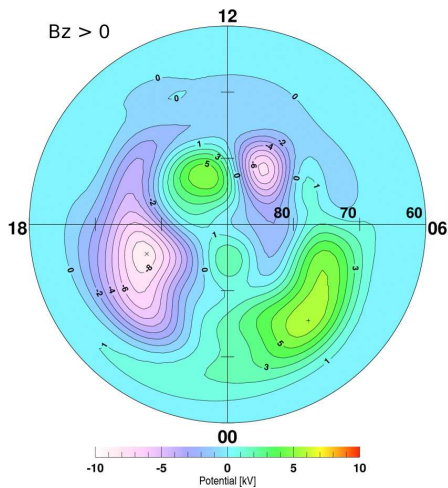
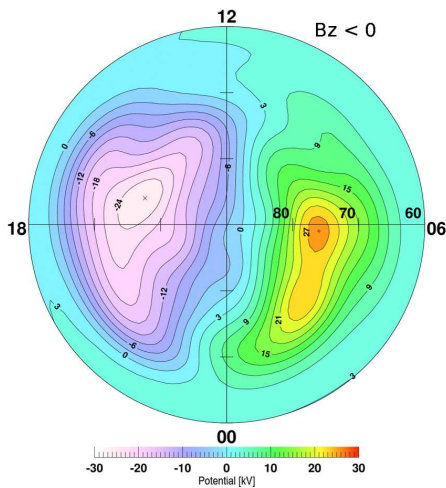


Milan and Grocott (2021)

Current Systems in the Ionosphere and Magnetosphere

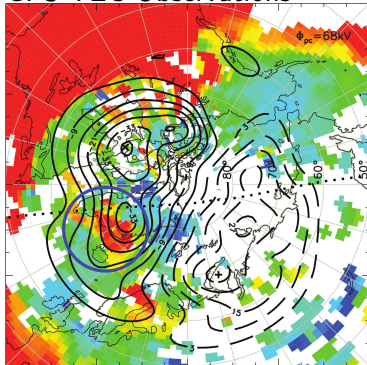


High Latitude Convection Patterns



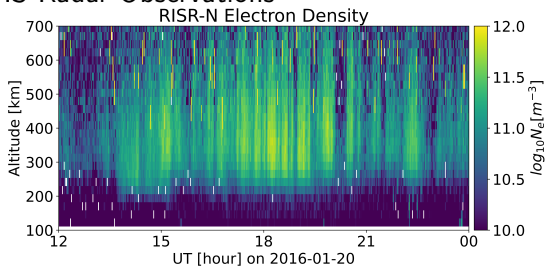
Density Structures in the Polar Cap

GPS TEC Observations

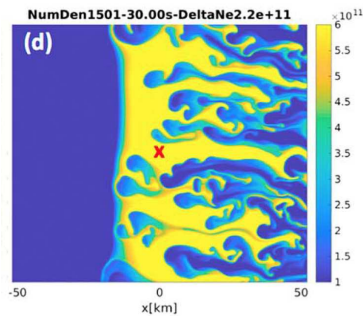
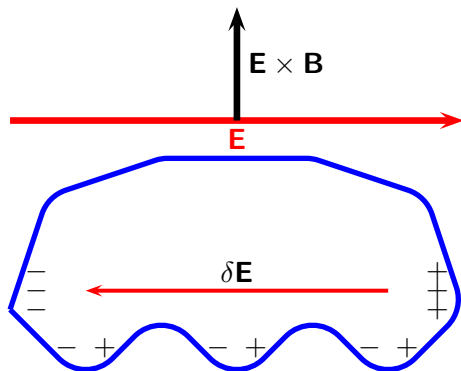


Zhang et al. (2013)

IS Radar Observations



Polarization Electric Fields and Gradient-Drift Instability



Deshpande and Zettergren
(2019) 10.1029/
2019GL082576

GDI Finger-like Structures

