

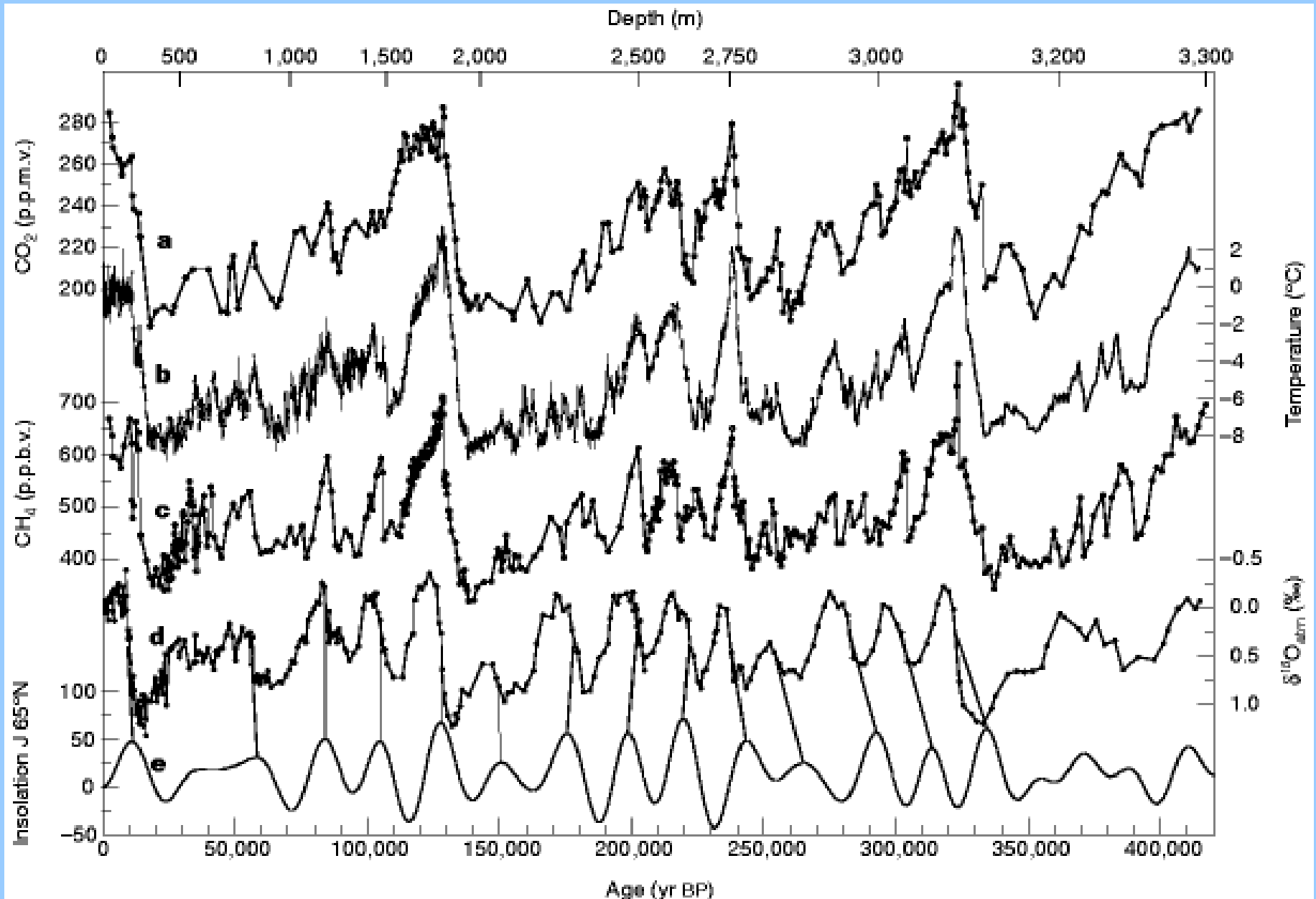
# A Very Brief Introduction to the Ocean Carbon Cycle

Mick Follows

Dec 2004

<http://ocean.mit.edu/~mick/Docs/carbon-intro.pdf>

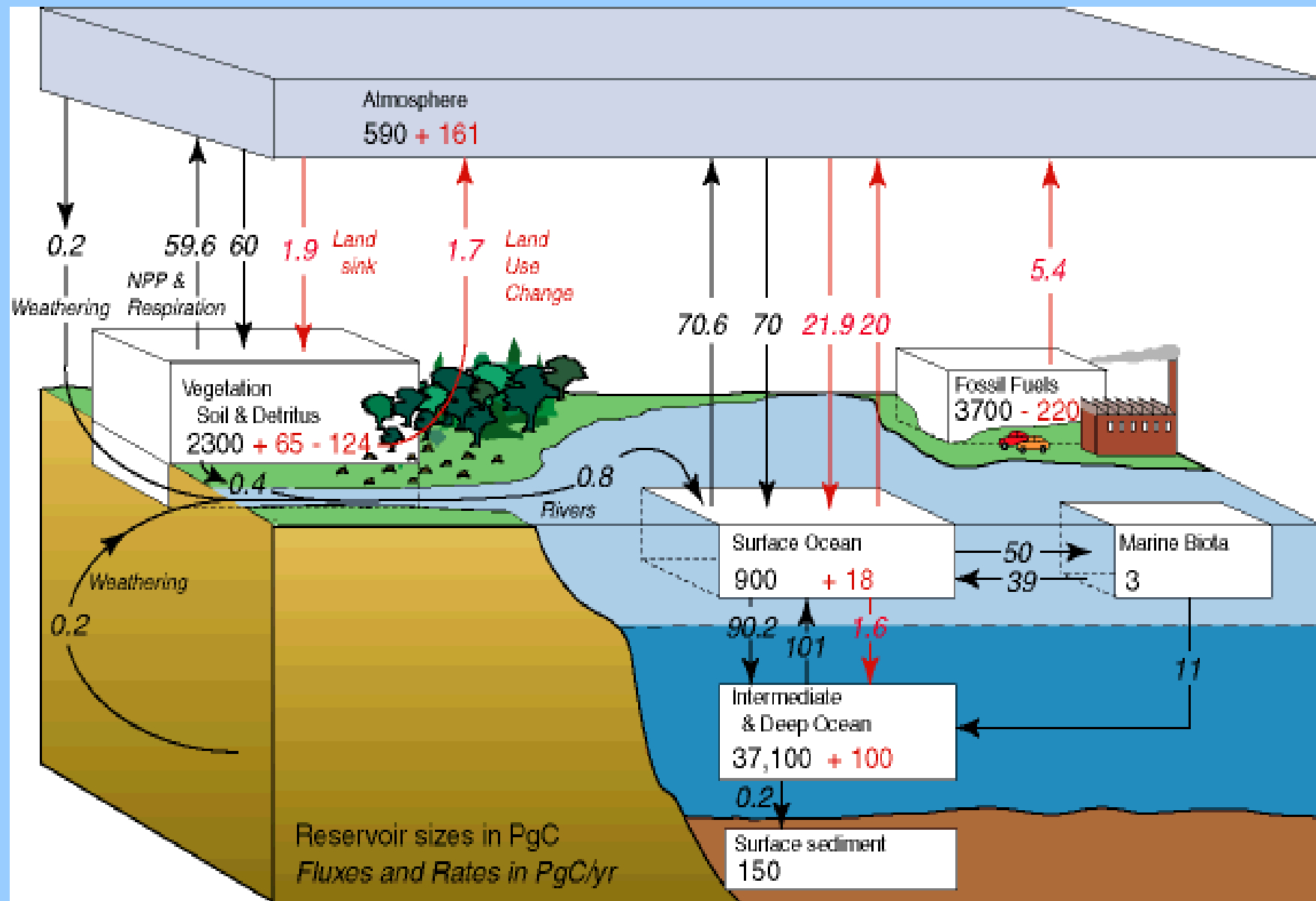
- What controls glacial-interglacial variation of atmospheric CO<sub>2</sub>?
- What is the connection to climate change?



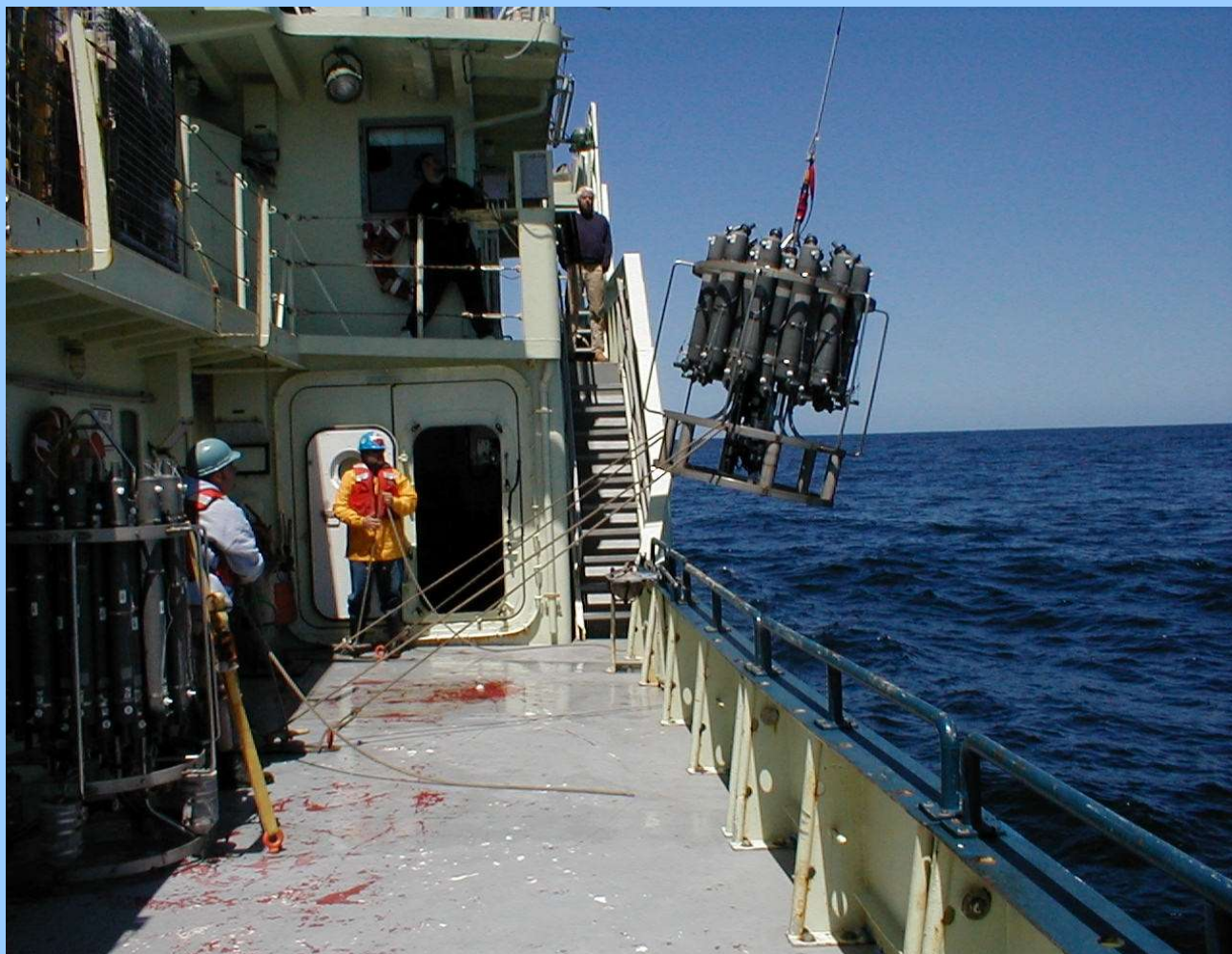
Ice core records from Vostock, Antarctica

(Petit et al., 1999)

# The Global Carbon Cycle

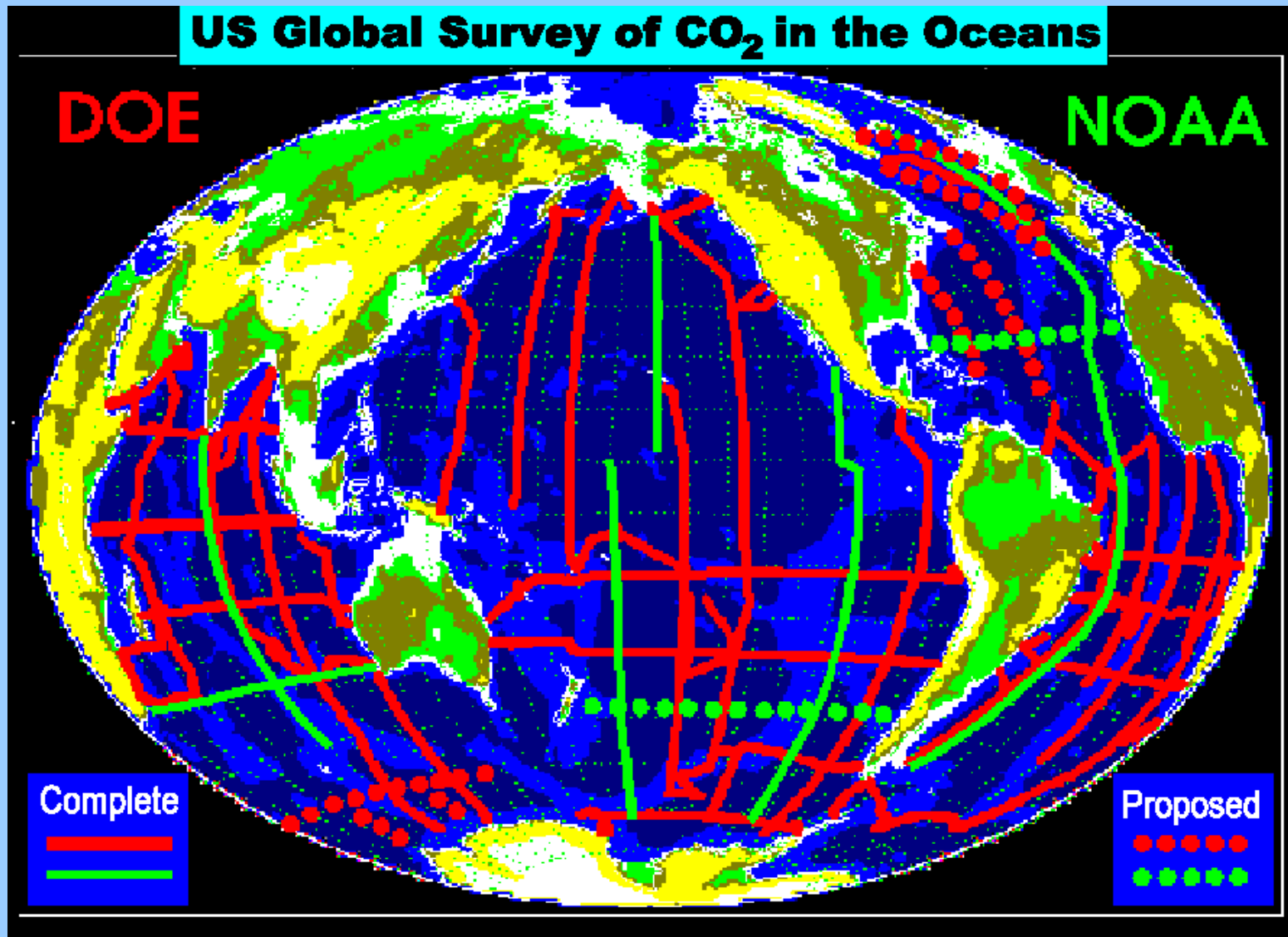


# Observing ocean biogeochemical properties



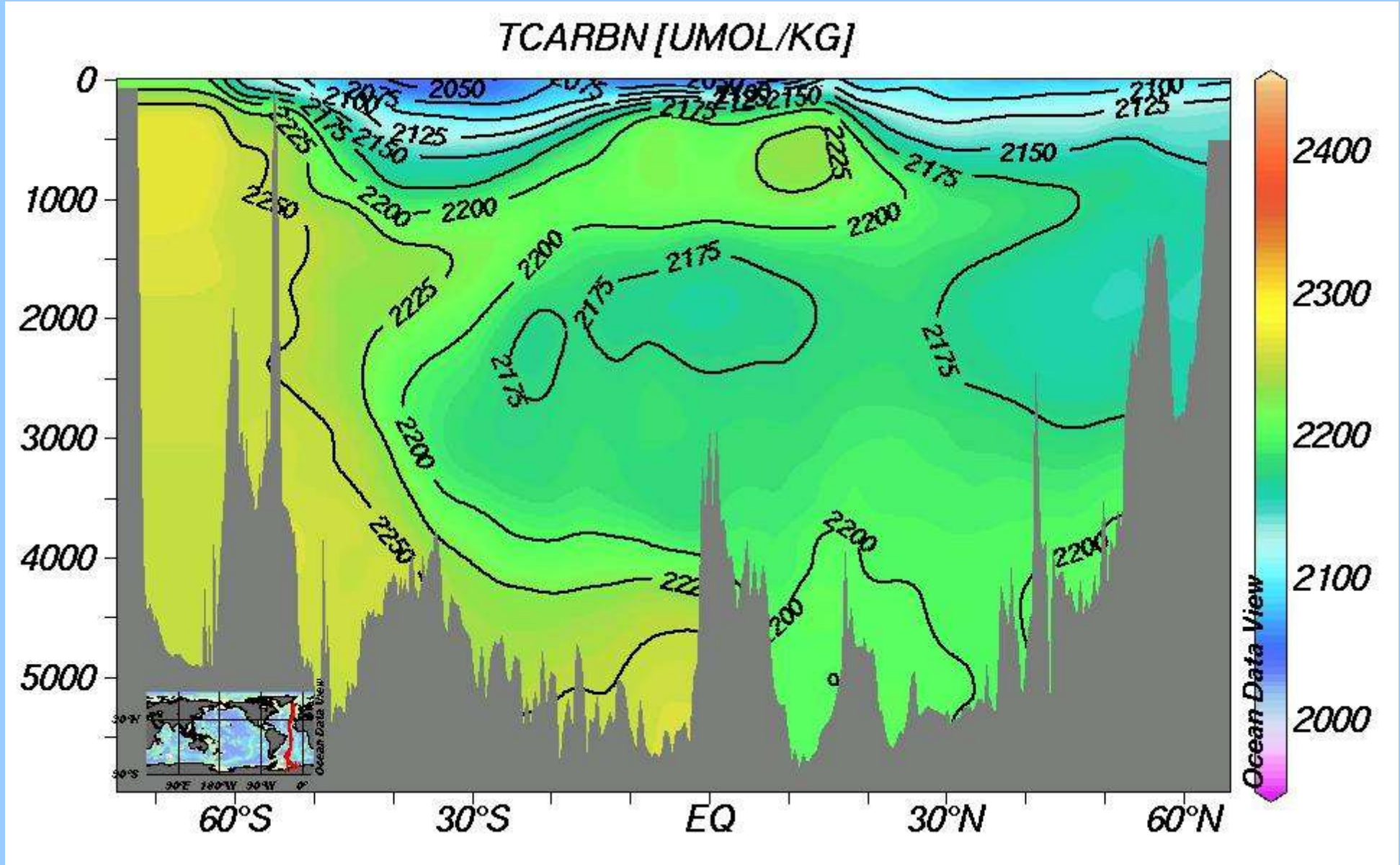
# What is the distribution of carbon in the ocean?

WOCE-JGOFS Global Ocean Survey (1990's)

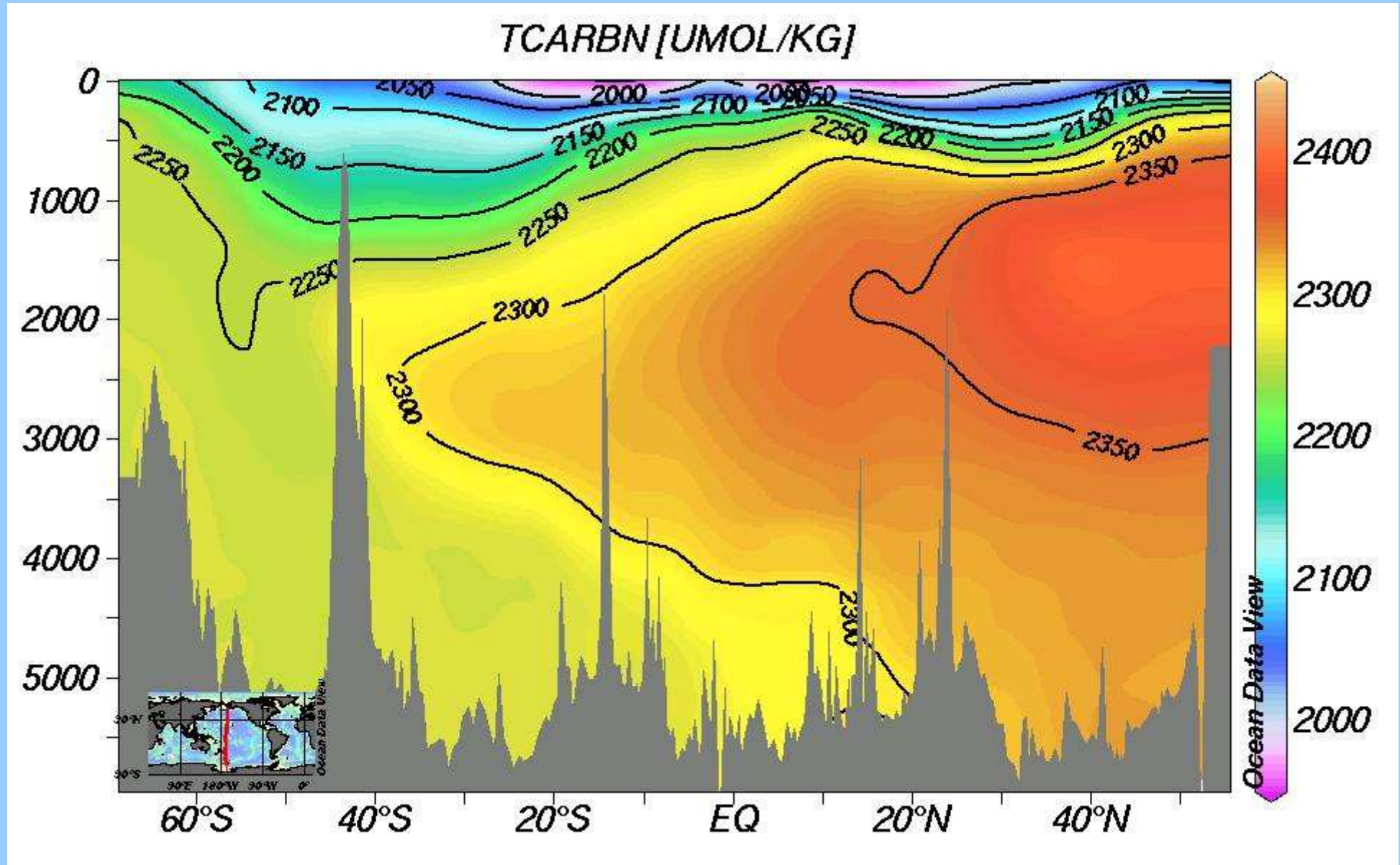




# Observed Atlantic Dissolved Inorganic Carbon, DIC

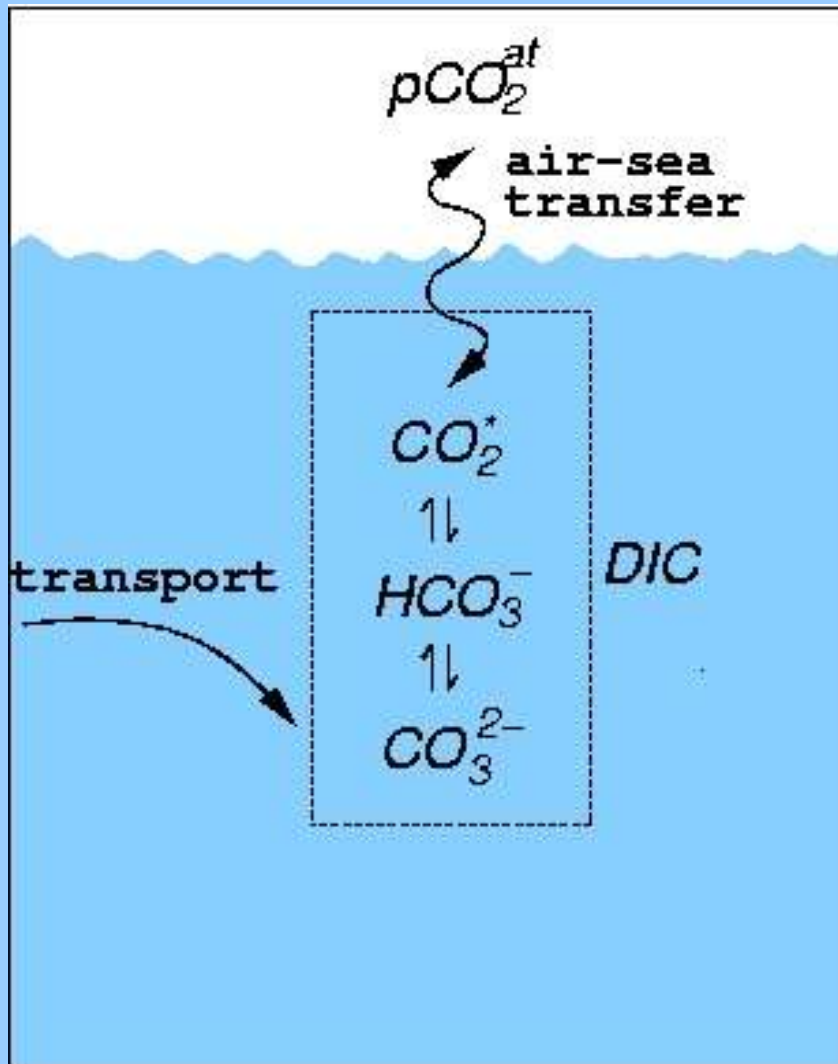


# Observed Pacific Dissolved Inorganic Carbon, DIC



- Background concentration of DIC  $\sim 2000 \text{ micromol kg}^{-1}$
- Why so much carbon in the ocean?

## Carbonate chemistry.



dissolved inorganic carbon

$$\text{DIC} = [\text{CO}_2^*] + [\text{HCO}_3^-] + [\text{CO}_3^{2-}]$$

$$\begin{array}{ccccc} [\text{CO}_2^*] & : & [\text{HCO}_3^-] & : & [\text{CO}_3^{2-}] \\ 1 & & 100 & & 10 \end{array}$$

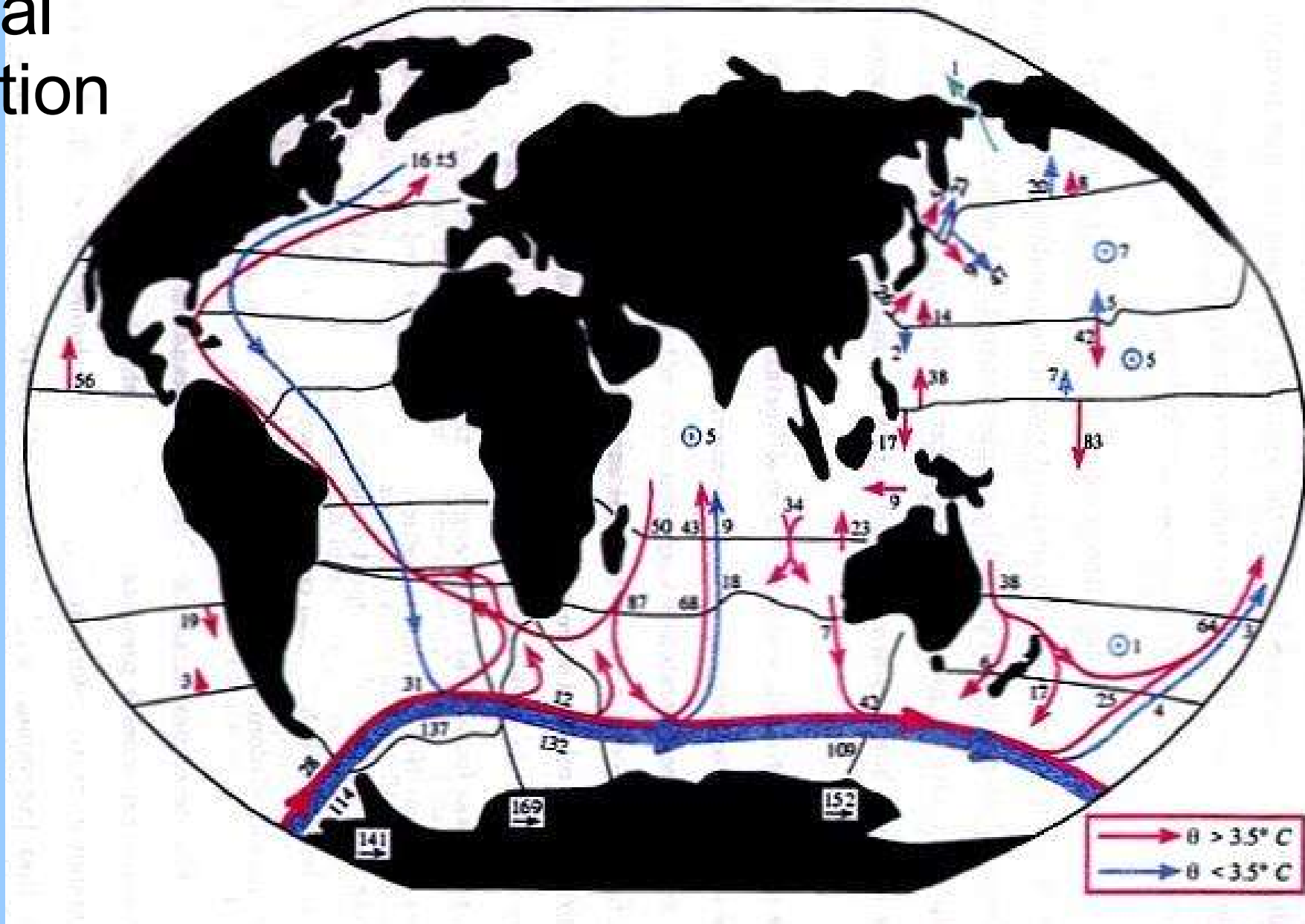


# What sets the distribution of carbon in the ocean?

- Combination of transport by circulation, chemical and biological processes.

1. Transport – relationship to S
2. “Solubility Pump” – relationship to T
3. “Biological Pumps” – relationship to  $\text{PO}_4^{3-}$

# Global circulation

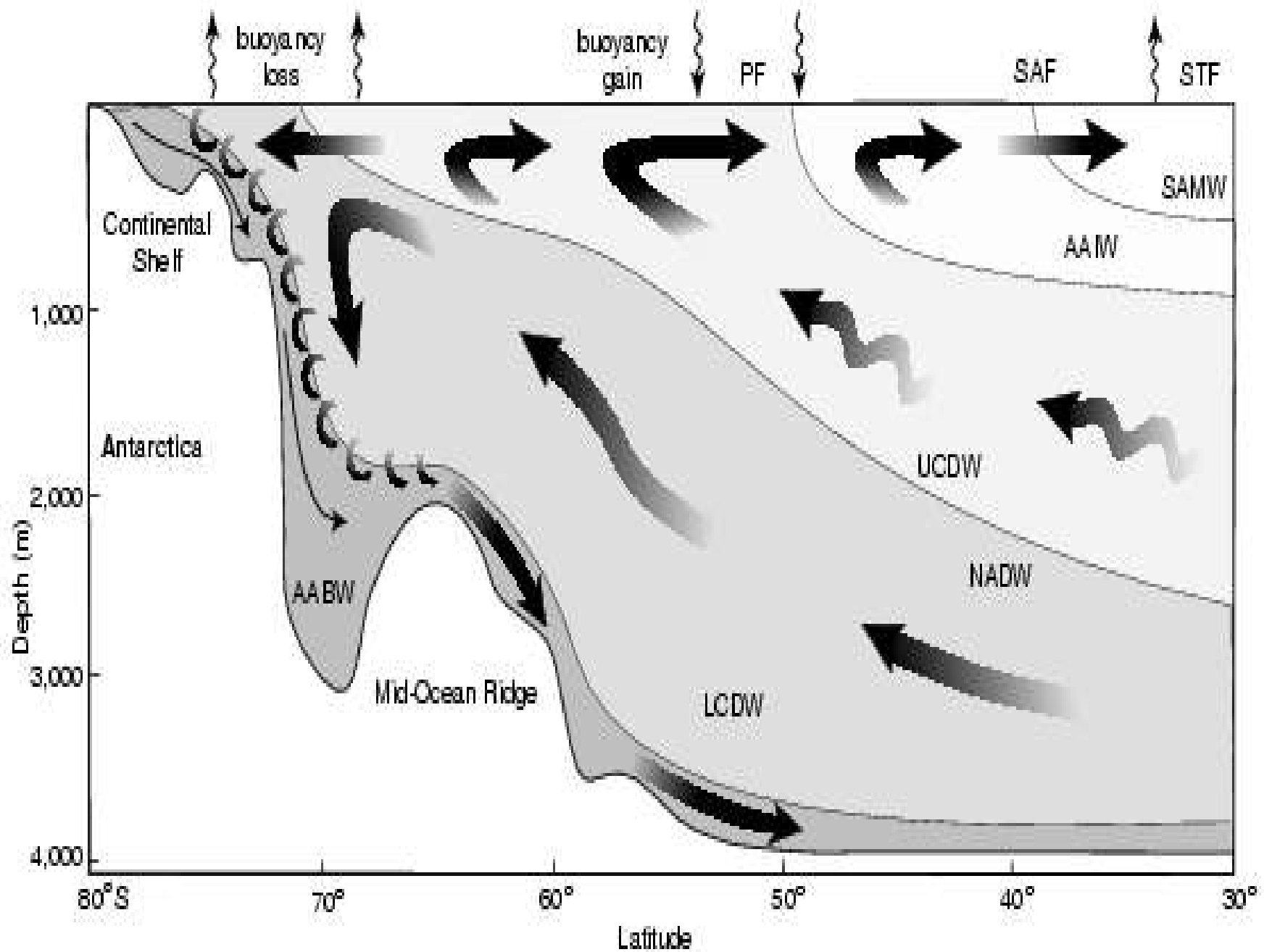


Volume flux of **warm** & **cold** water

*Macdonald and Wunsch (1996) inversion*

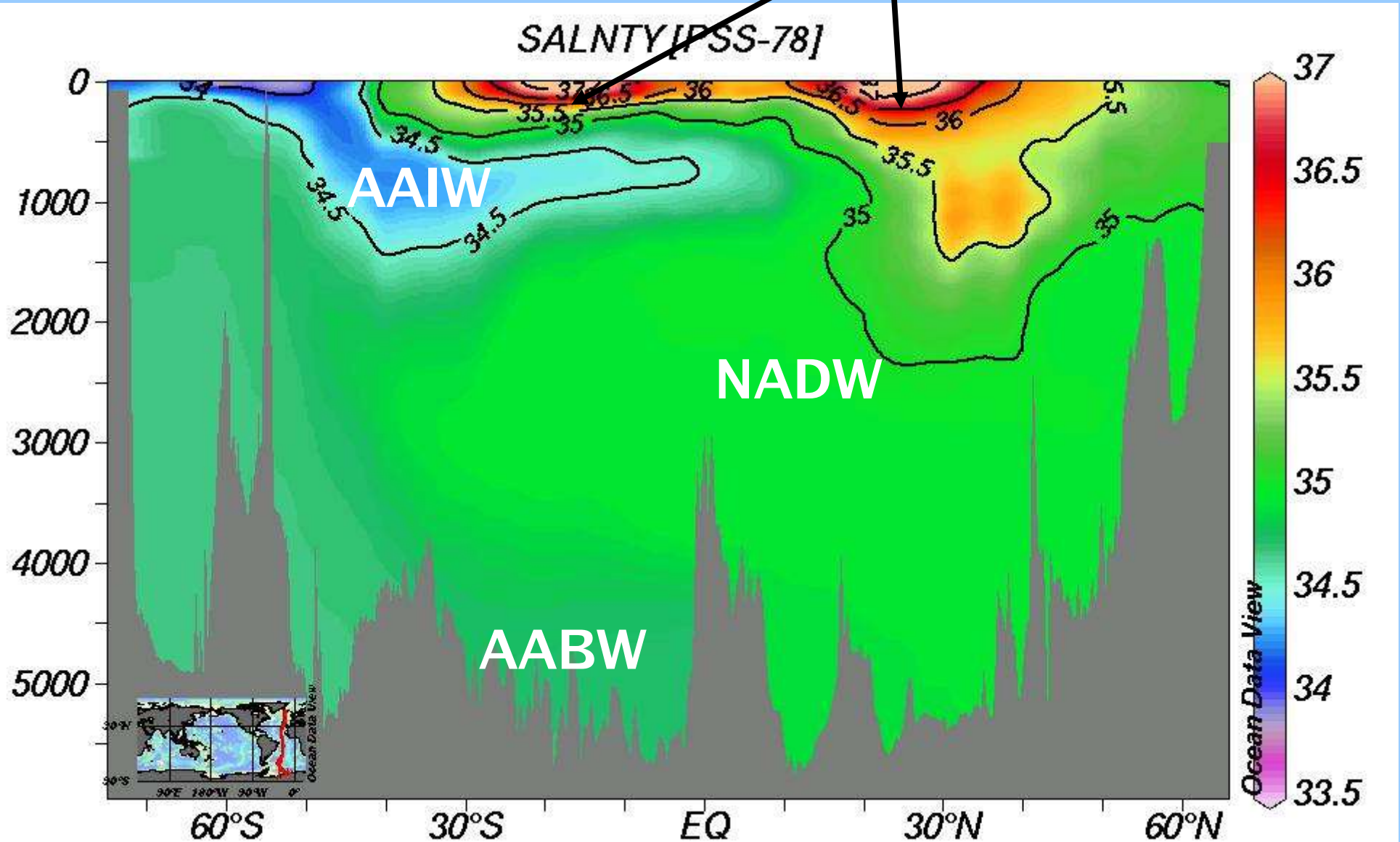
# Southern Ocean

*Speer et al. (2000)*



Atlantic, salinity (psu)

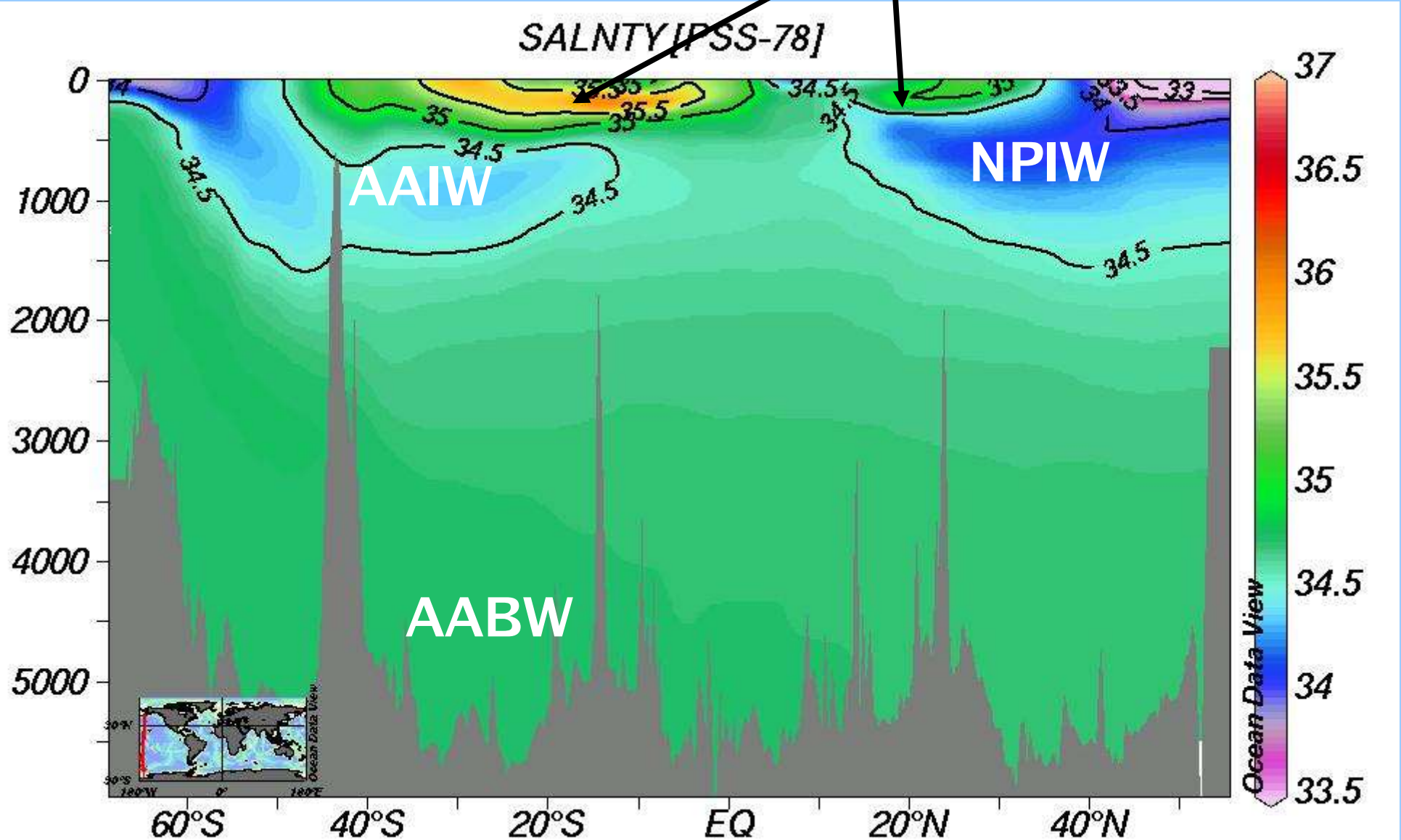
ventilated thermocline

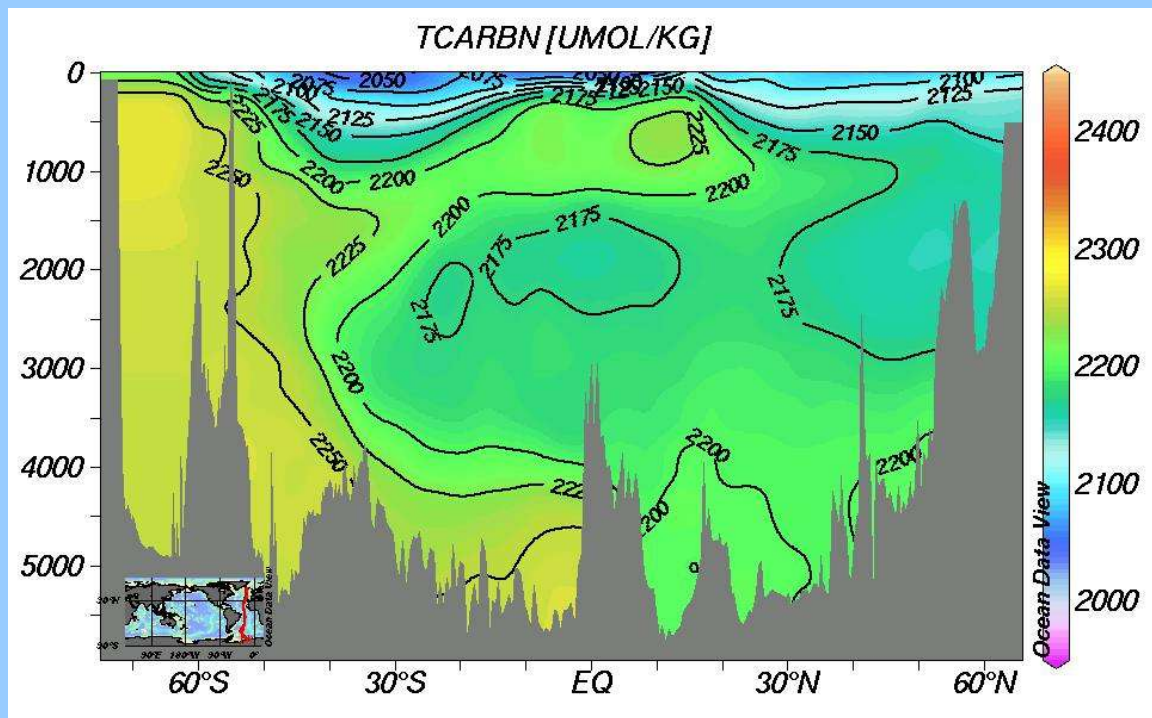




Pacific, salinity (psu)

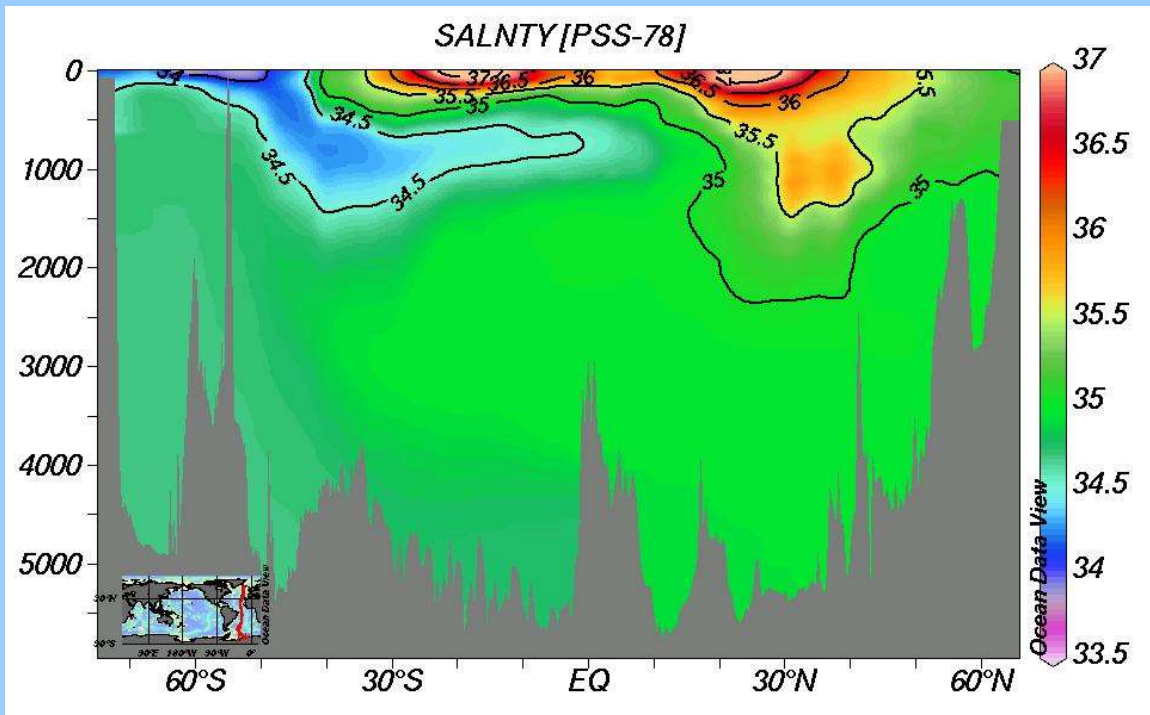
ventilated thermocline

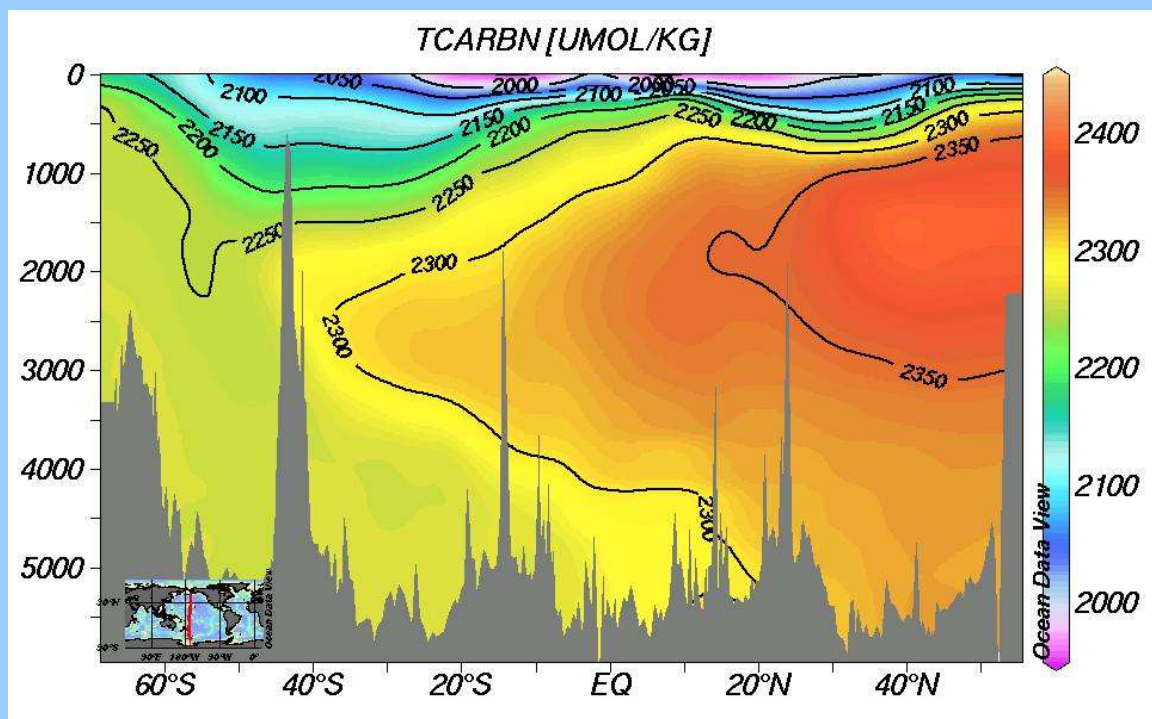




DIC and S distributions show some common structures:

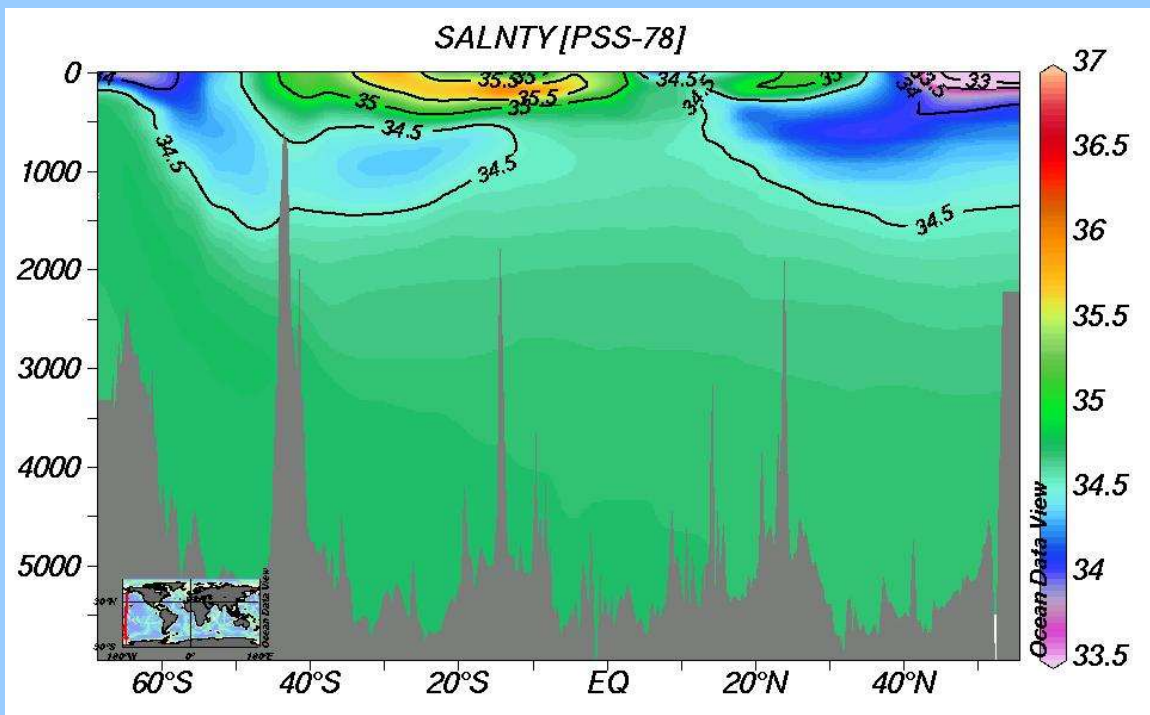
- Common ocean transport processes





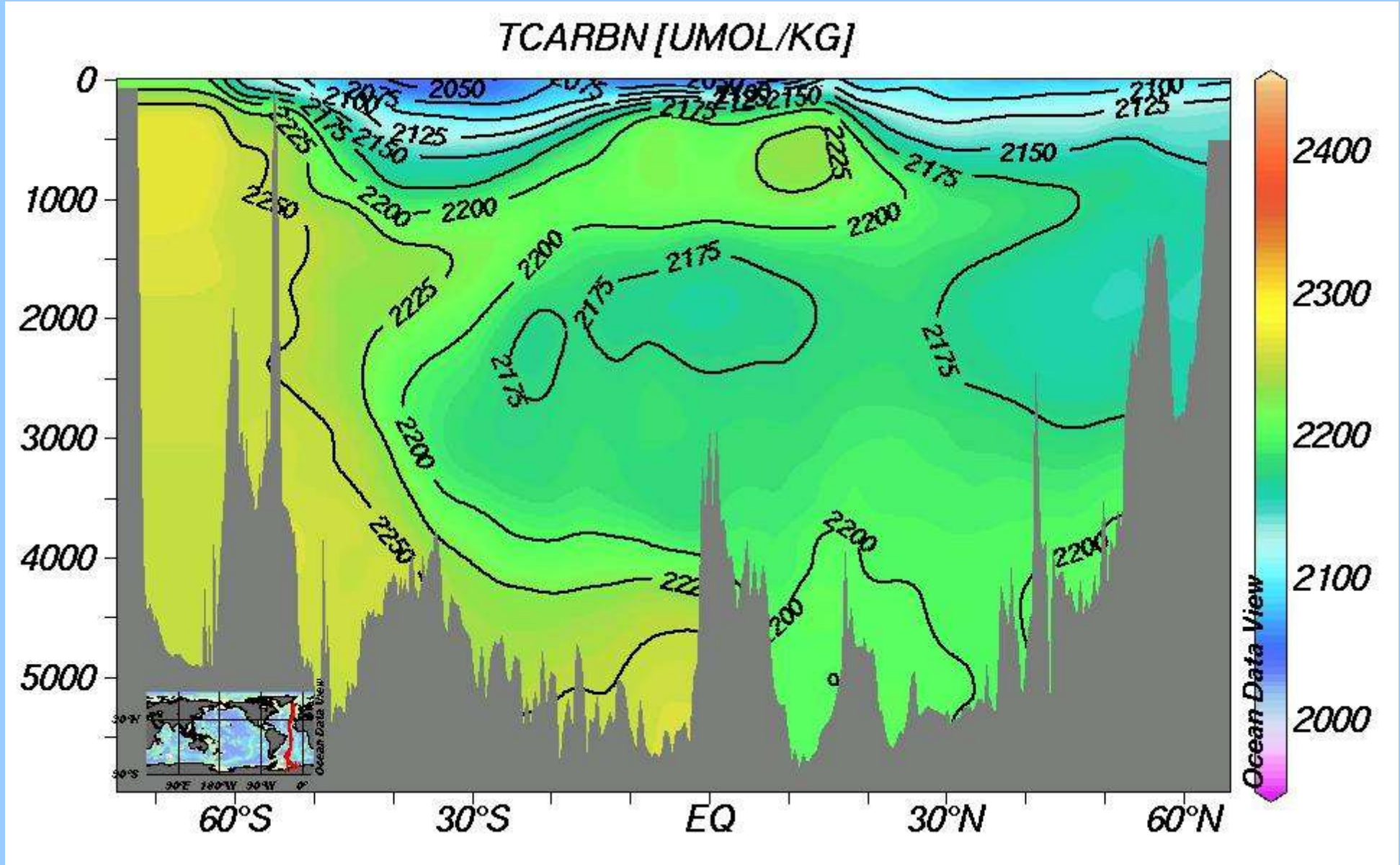
DIC and S distributions show some strong differences:

- DIC affected by biological processes but S is not



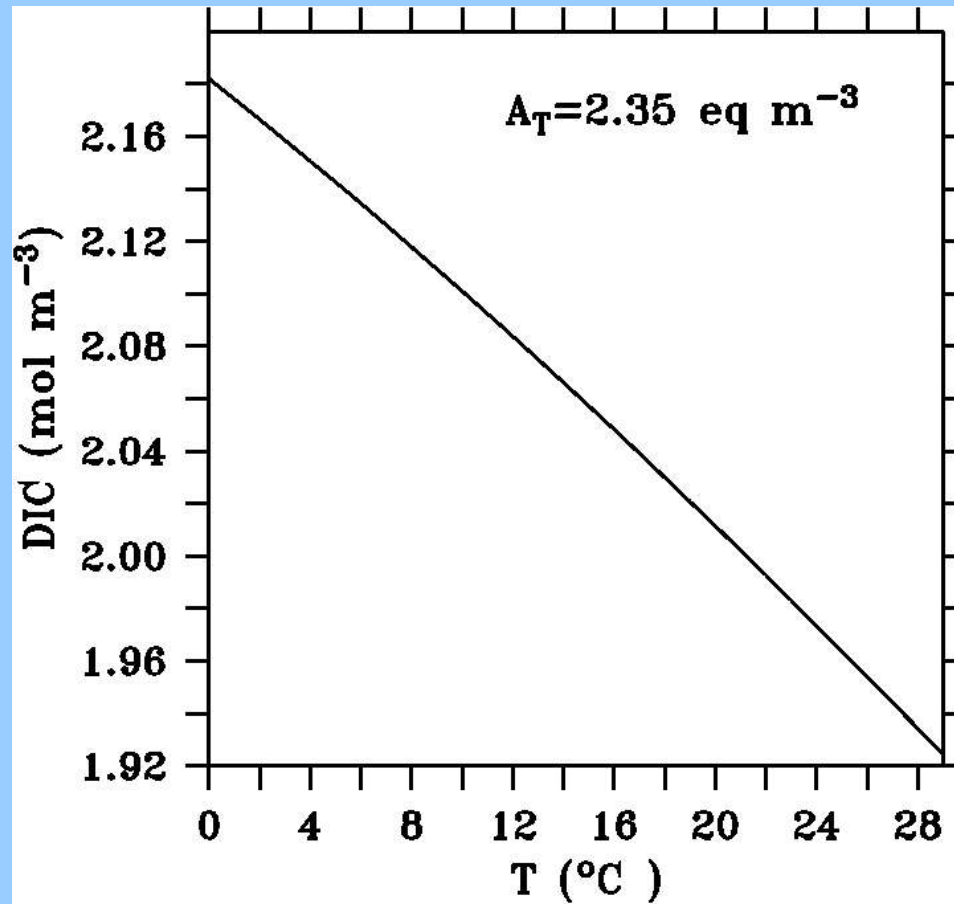


# Observed Atlantic Dissolved Inorganic Carbon, DIC





# Relationship of DIC at equilibrium with overlying atmos $p\text{CO}_2=280\text{ppmv}$

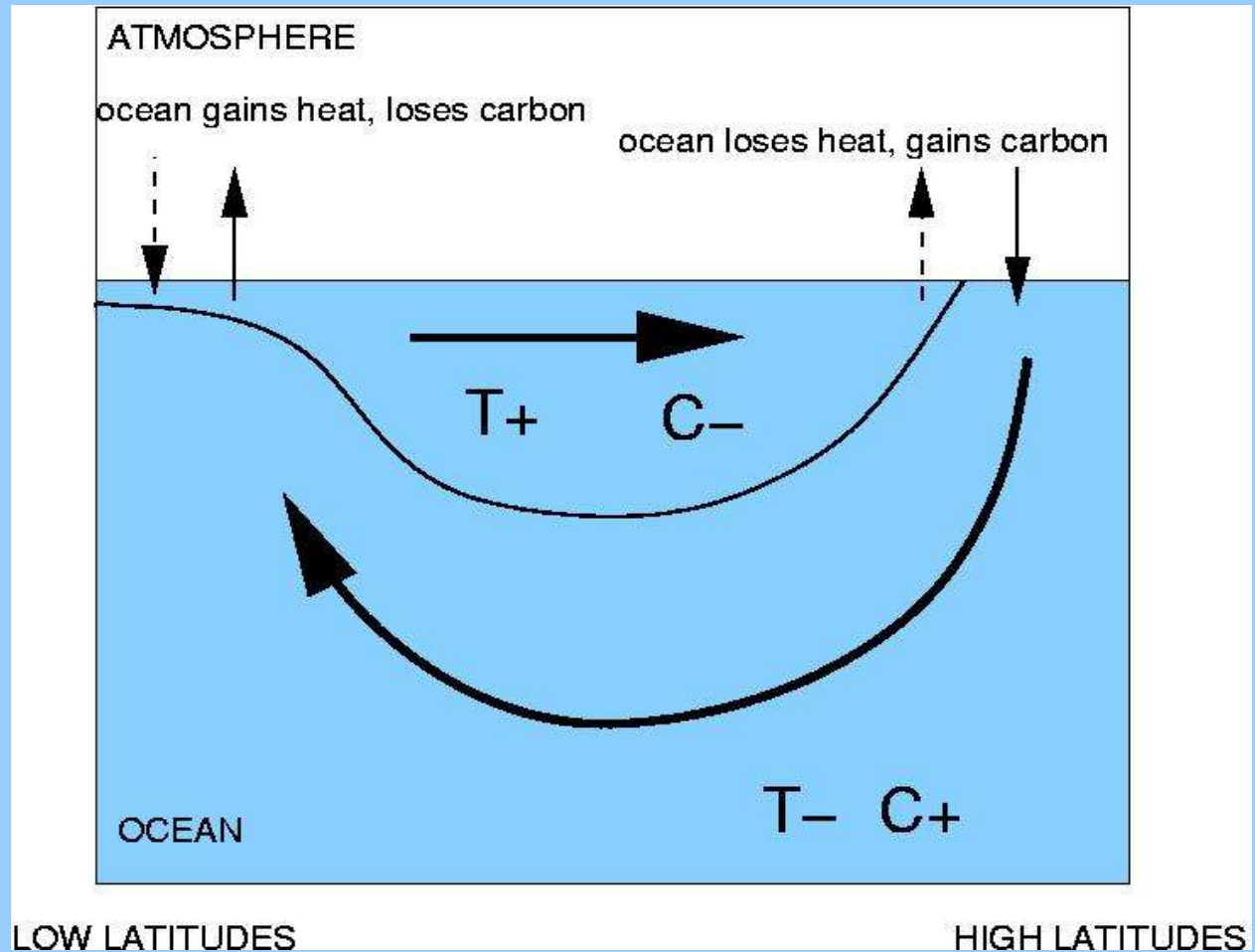


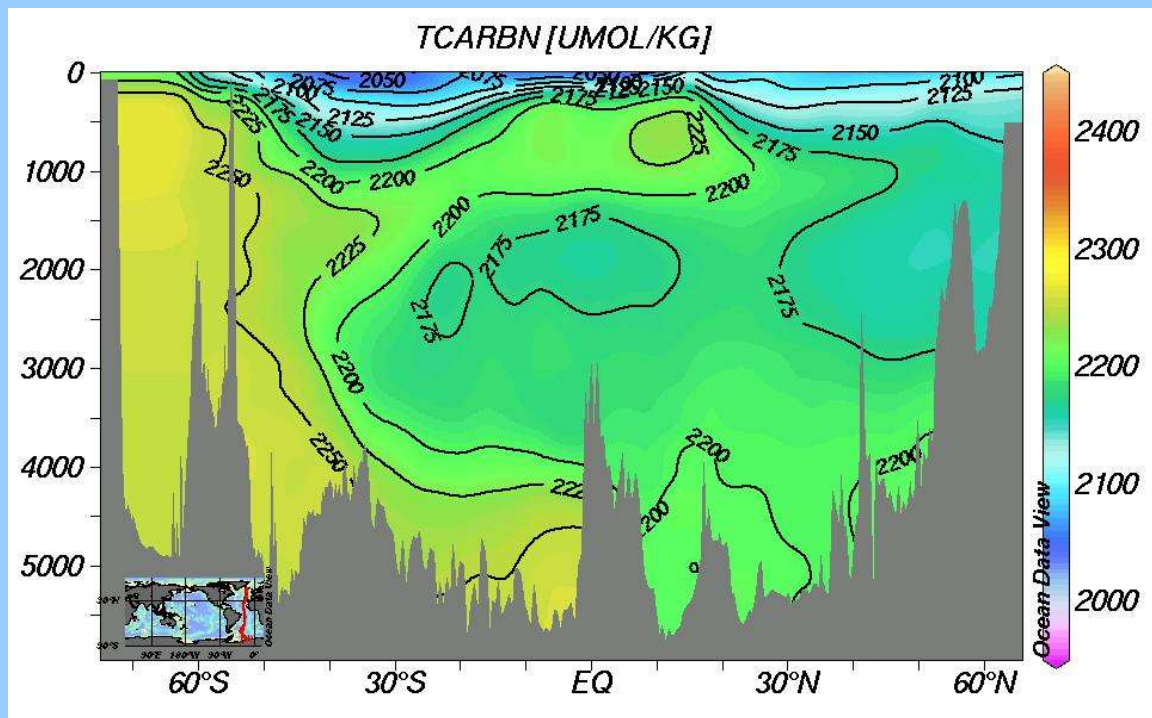
**Cooler waters hold more carbon at equilibrium**

$$\text{mol m}^{-3} = \text{micromol kg}^{-1} / 10^3$$

# “Solubility Pump”

- cooler water holds more DIC at equilibrium for given atmos  $p\text{CO}_2$
- cooler waters are denser and form deep waters
- creates vertical gradient of DIC in the ocean
- air-sea heat fluxes drive air-sea carbon fluxes

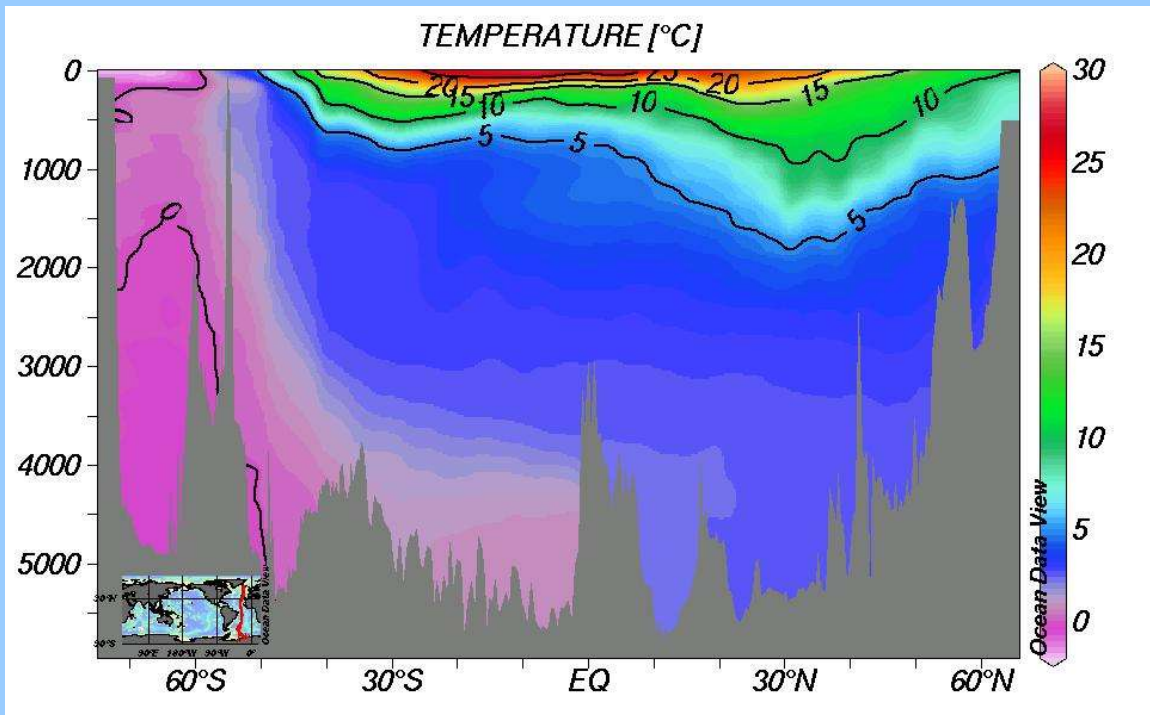


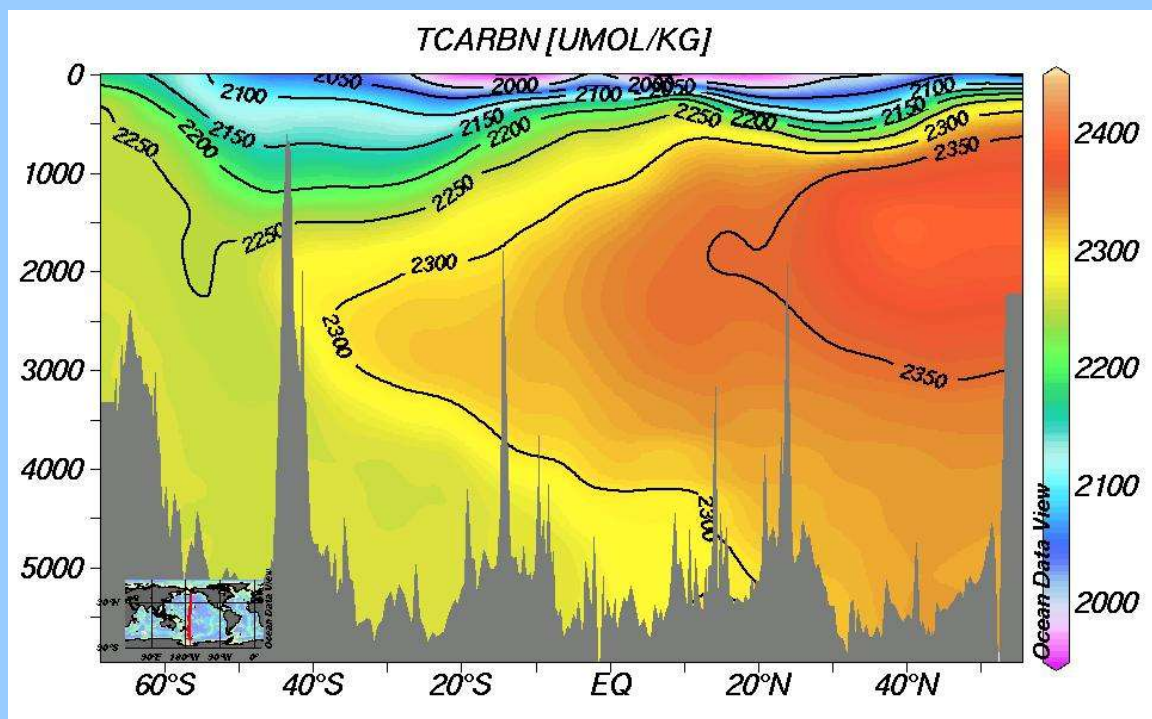


## Atlantic

Vertical gradients of  
DIC and T generally  
inversely related:

## Solubility Pump

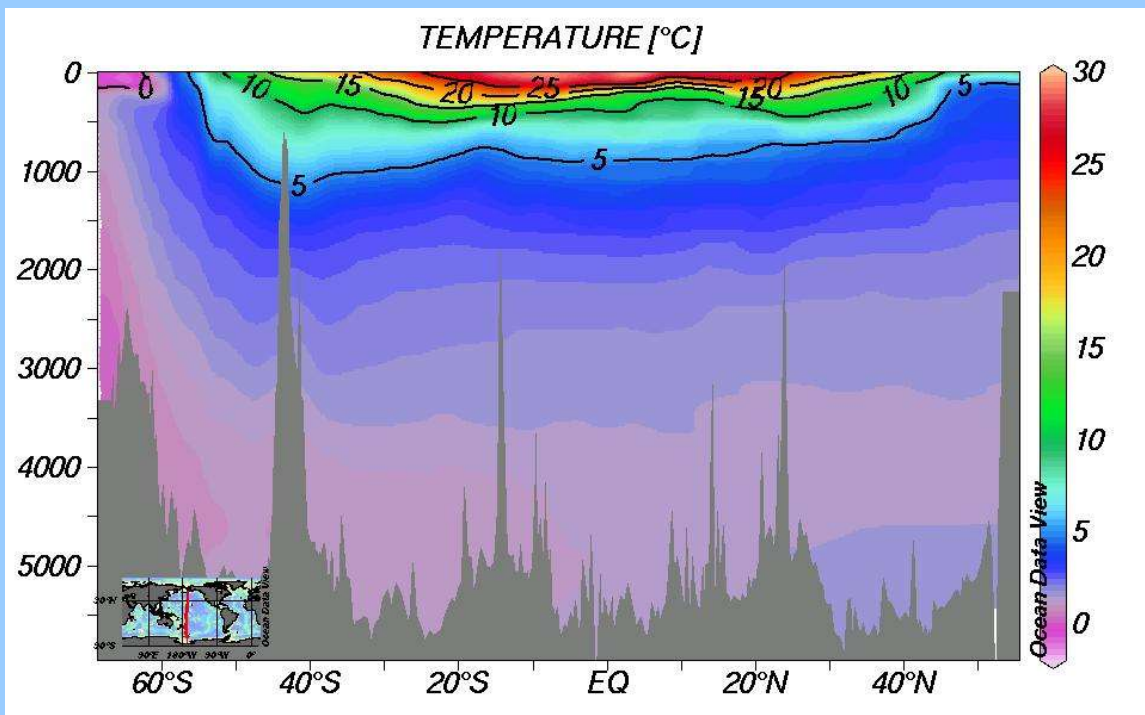




## *Pacific*

Vertical gradients of  
DIC and T generally  
inversely related

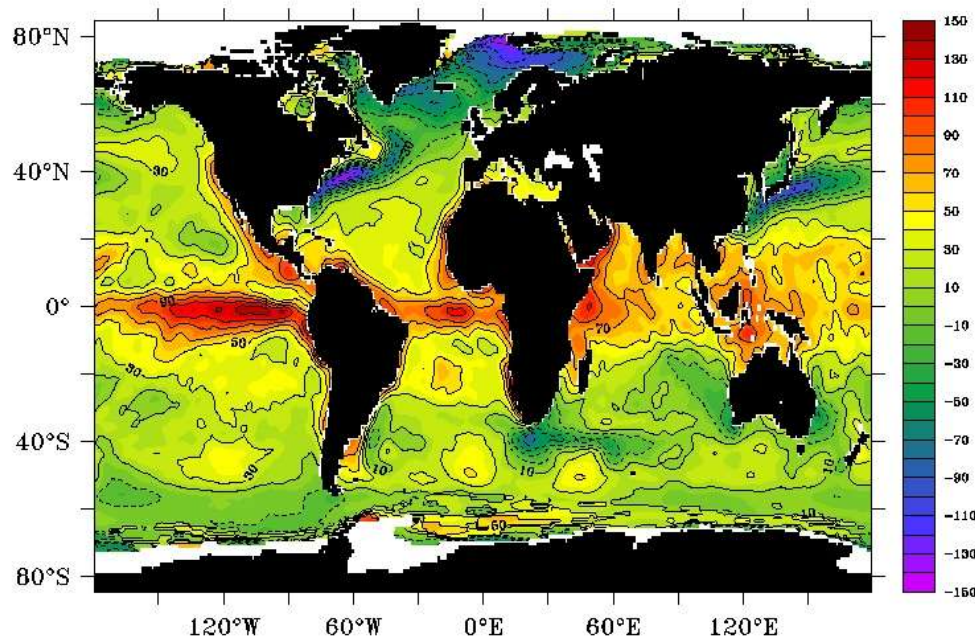
**Solubility pump**





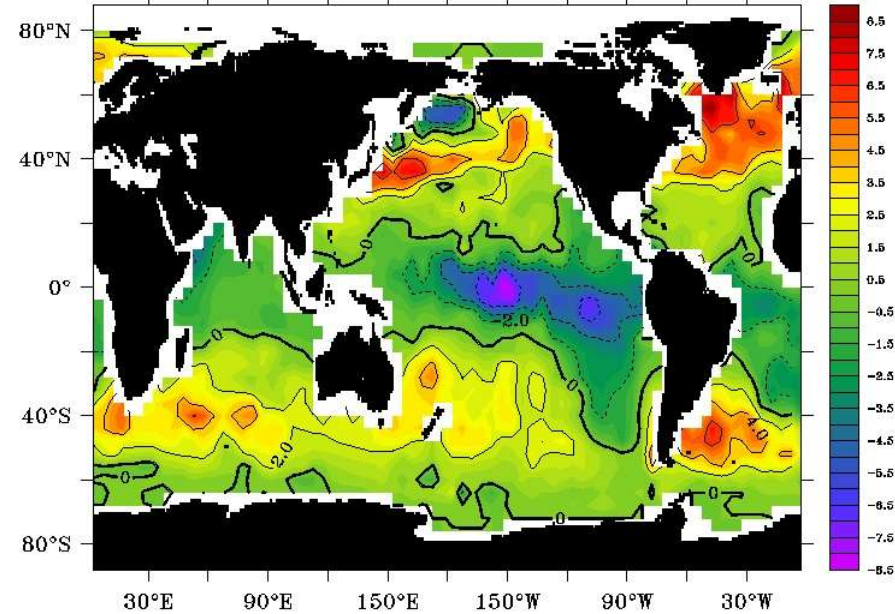
# Air-sea flux of CO<sub>2</sub> inversely related to air-sea heat flux

SOC net surf heat flux ( $\text{W m}^{-2}$ )



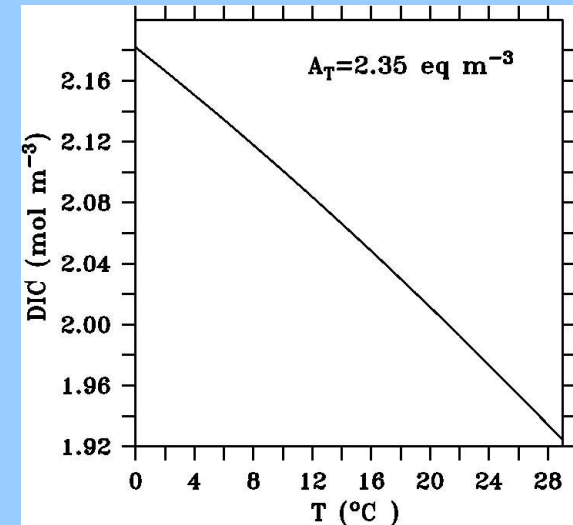
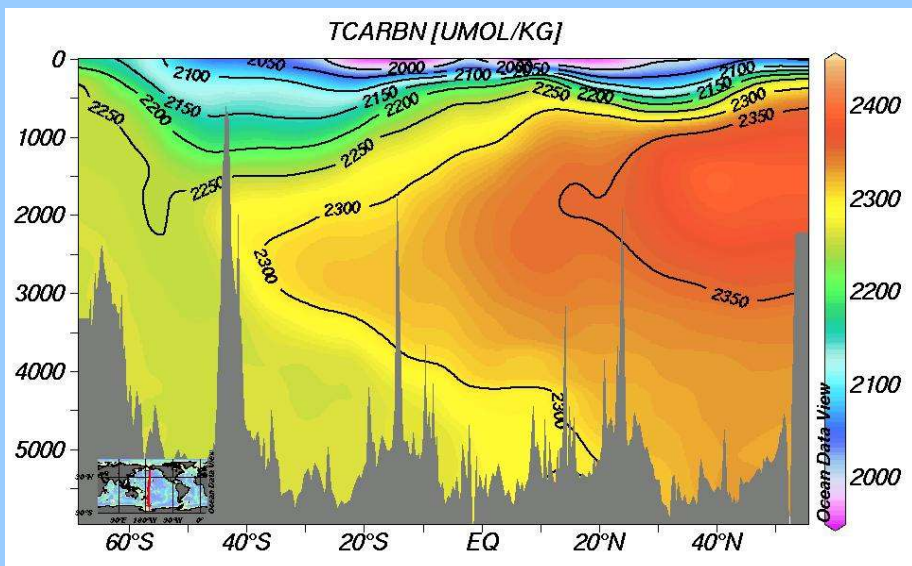
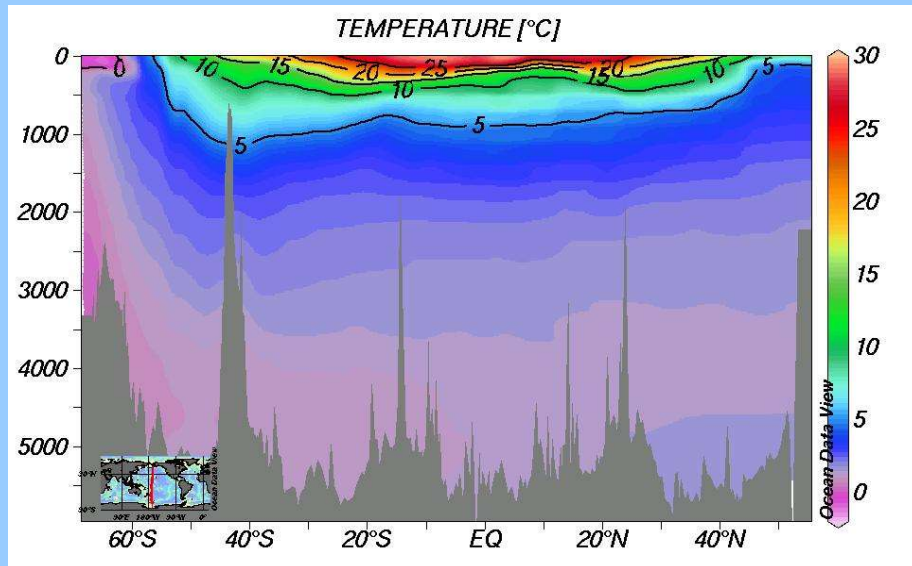
red – ocean gaining heat

Takahashi et al. (1997) Air-sea flux of CO<sub>2</sub> ( $\text{mol m}^{-2} \text{ yr}^{-1}$ )



red – ocean gaining carbon

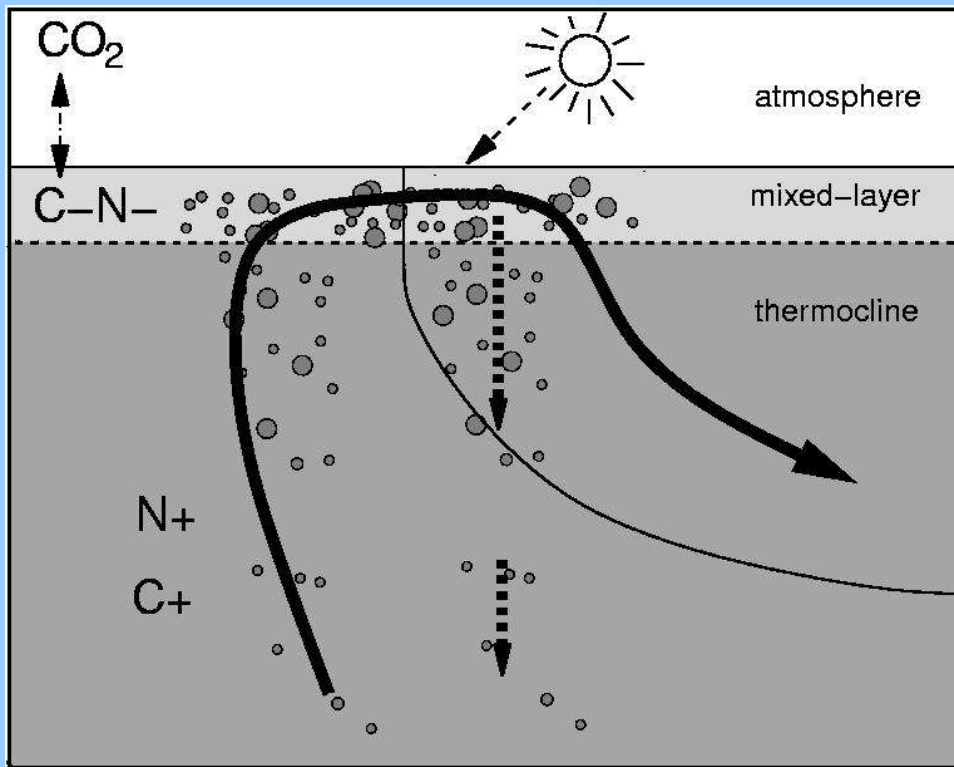
But DIC in the deep ocean exceeds the equilibrium concentration at that temperature... why?



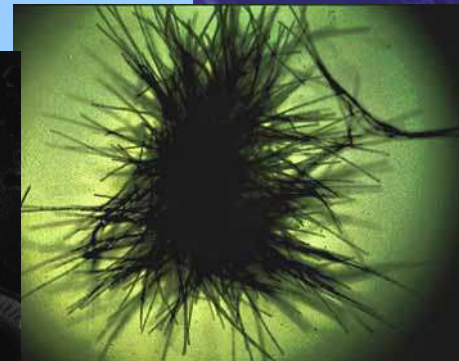
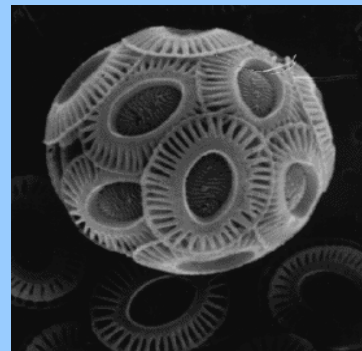
# Biological (soft tissue) Pump

**Photosynthesis** occurs in sunlight region (upper 150m)  
Requires light and nutrients, creates oxygen

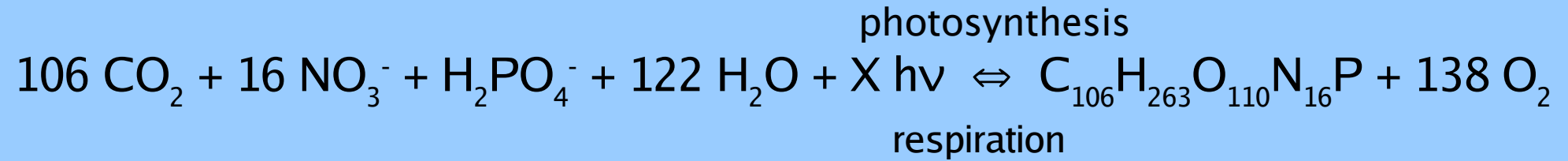
**Respiration**; remineralization of organic matter, consumes oxygen



Phytoplankton

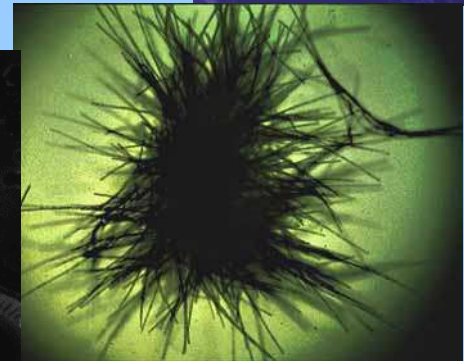
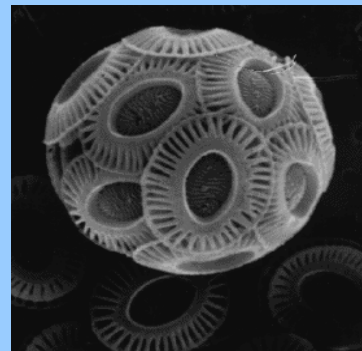
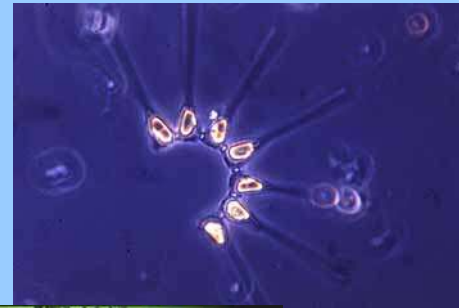


# Biological (soft tissue) Pump



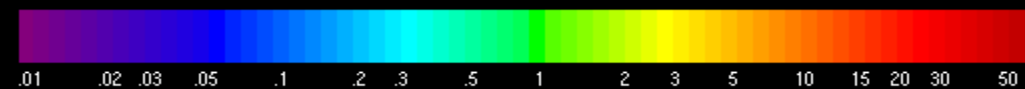
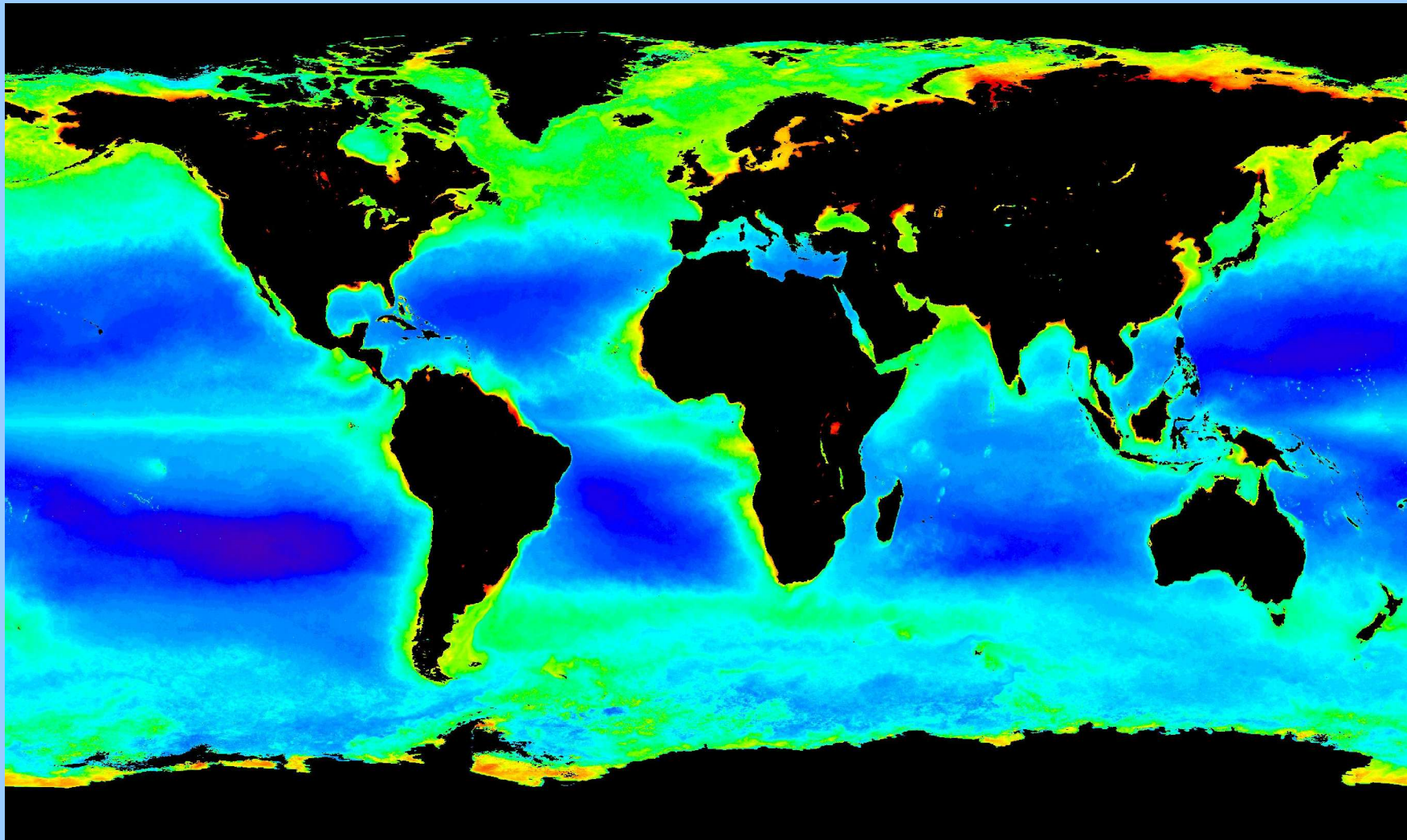
**Redfield Ratio** – oceanic organic matter has “fixed” ratio of carbon and nutrients

Phytoplankton





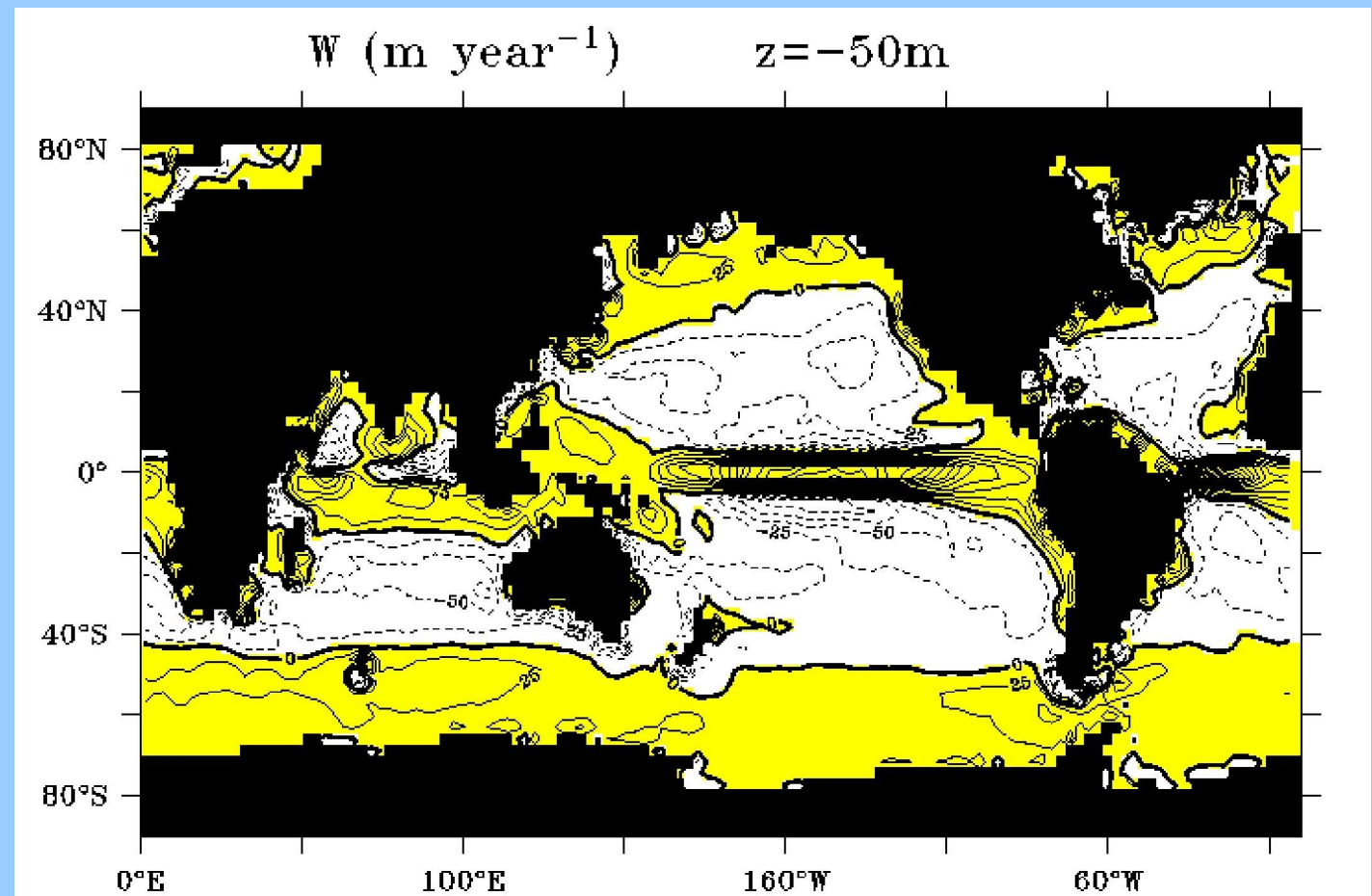
# Distribution of chlorophyll in surface ocean observed from space (SeaWiFS, 1997-2000 mean)



Chlorophyll a Concentration  
mg/m3

Lots of carbon around so productivity is limited by availability of other nutrients (e.g.  $\text{PO}_4^{3-}$ ) and light

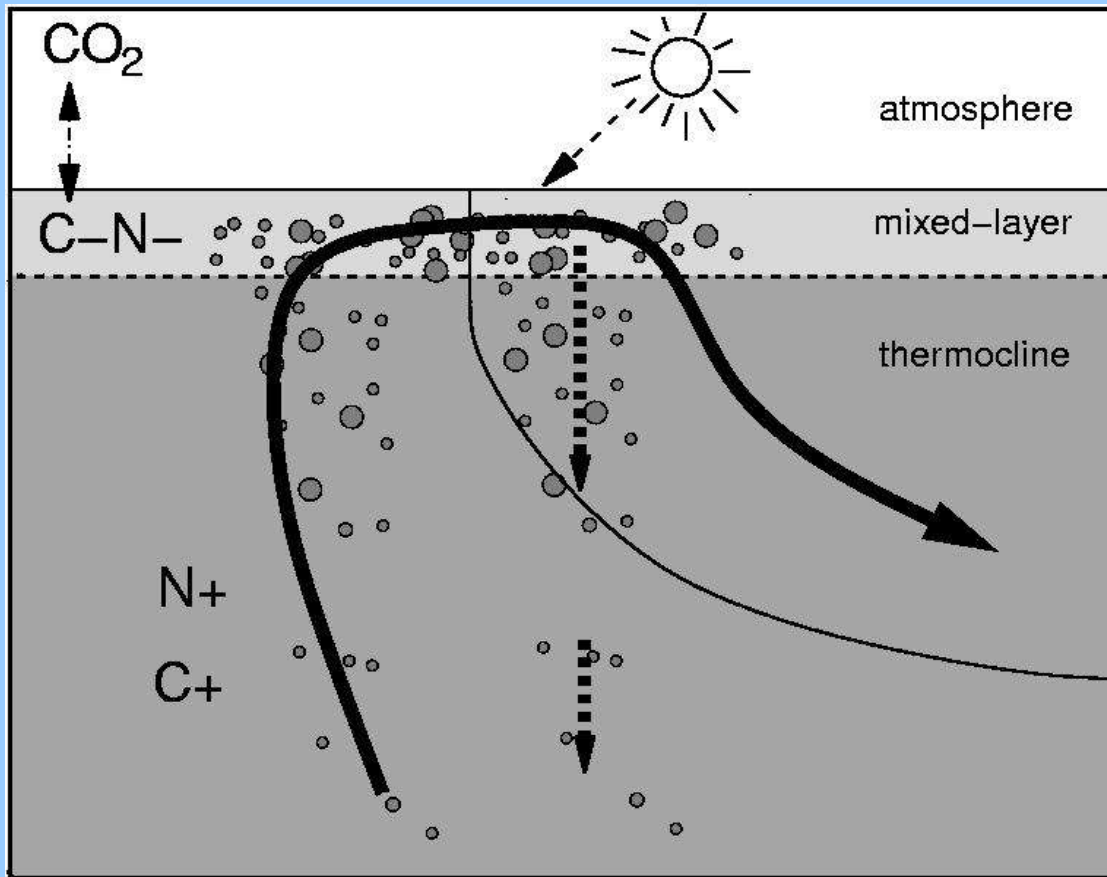
Supply of nutrients to sunlit surface ocean largely dependent on wind driven upwelling



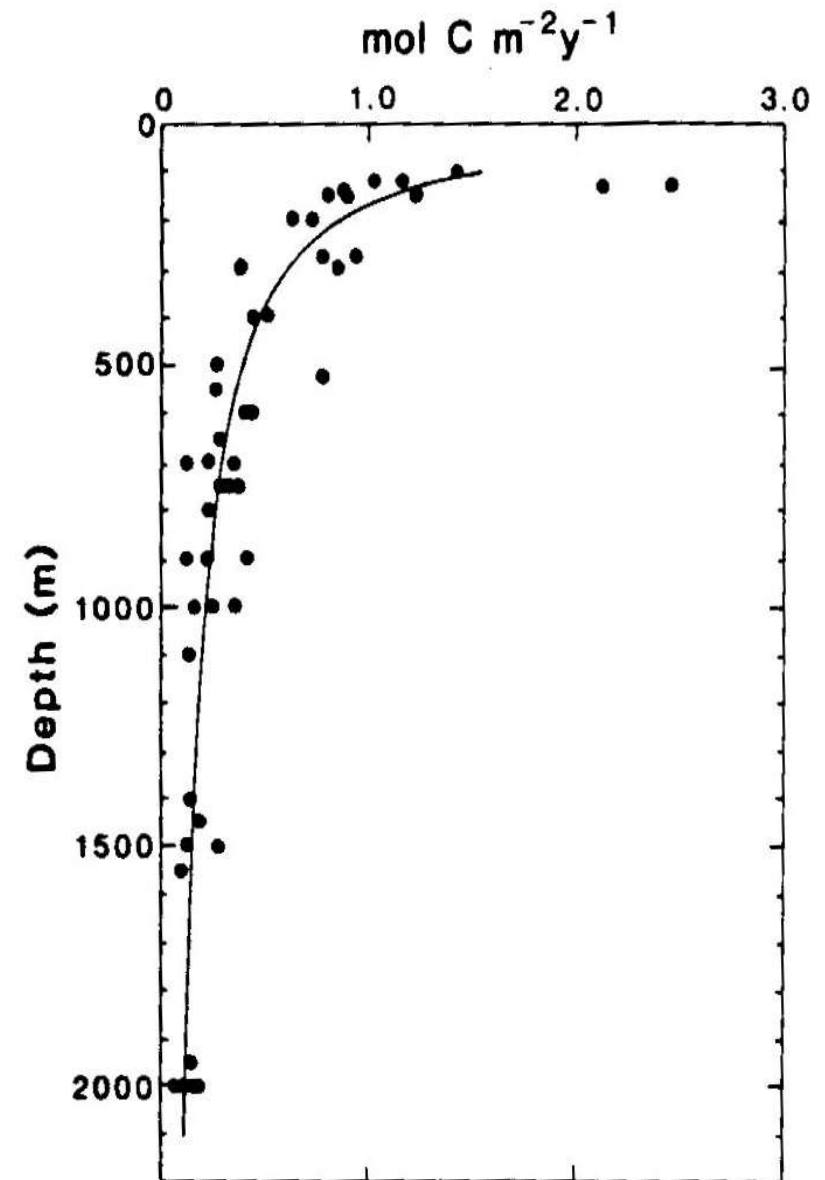
Modelled upwelling at 50m

# Organic carbon flux in sinking particles

Martin et al. (1987)



“Soft tissue pump”

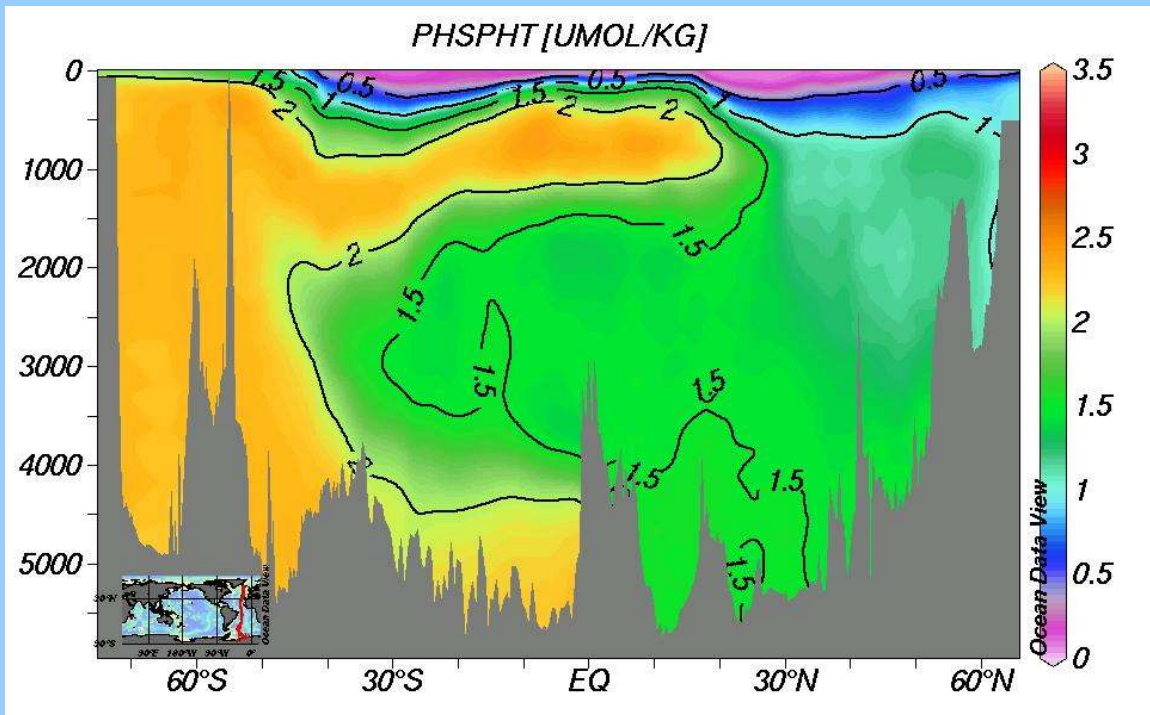
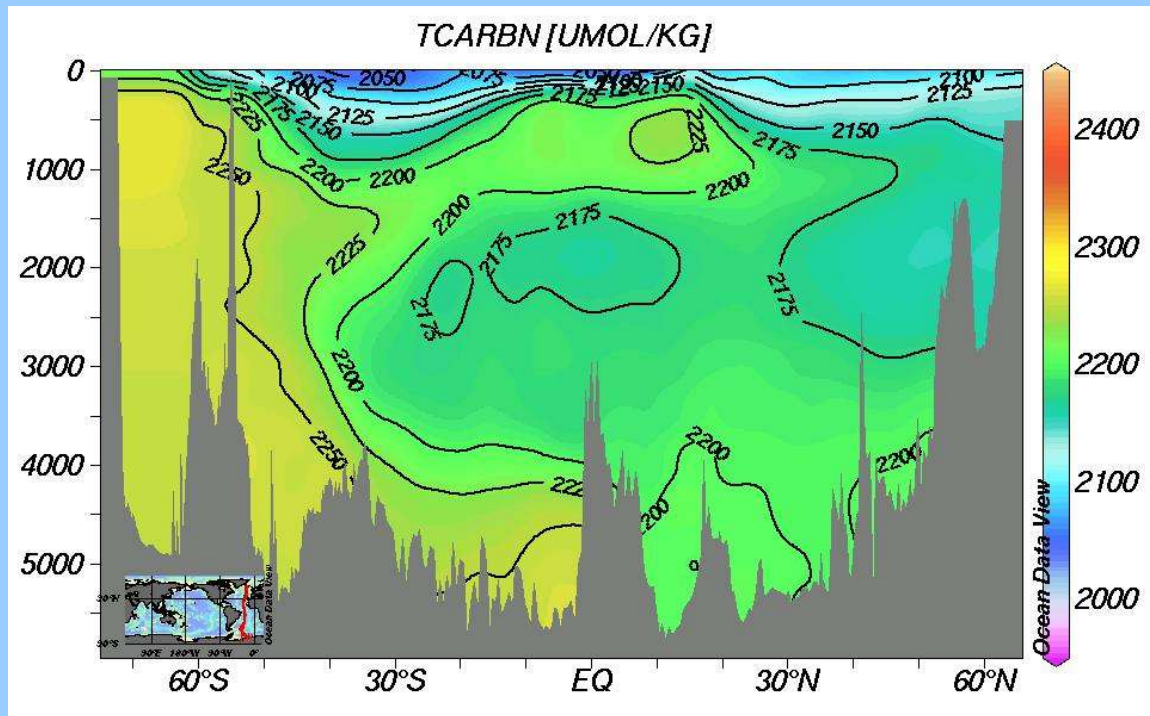


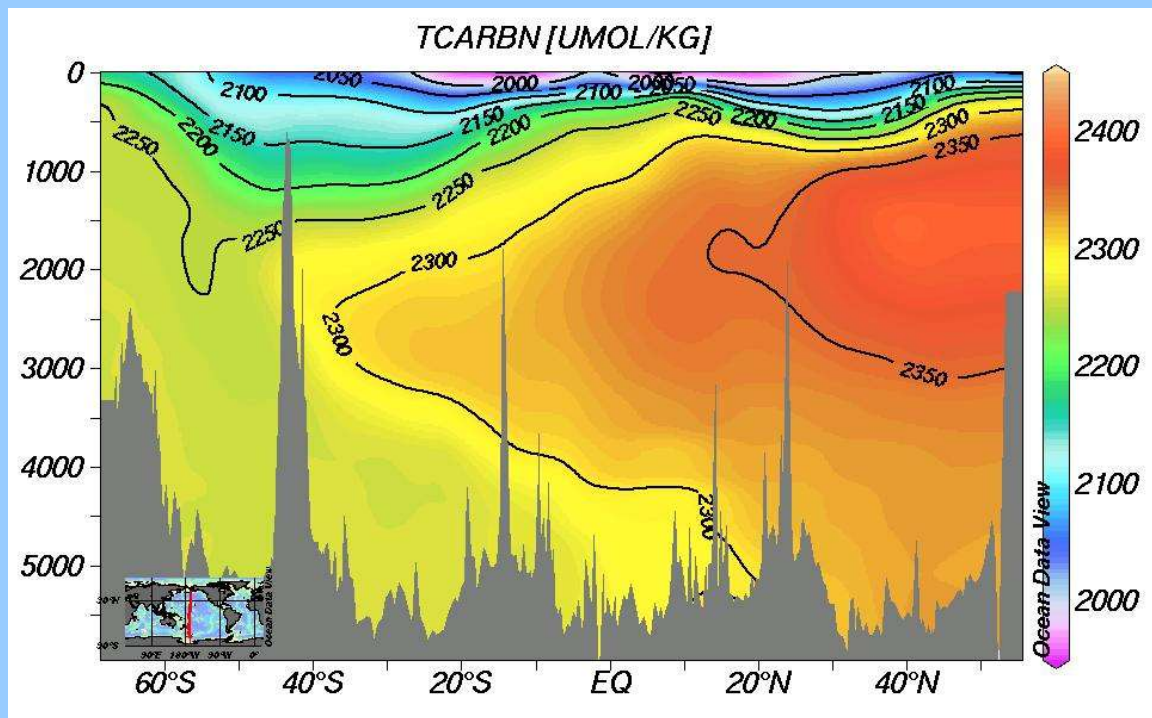


## DIC and PO4

- common biological processes

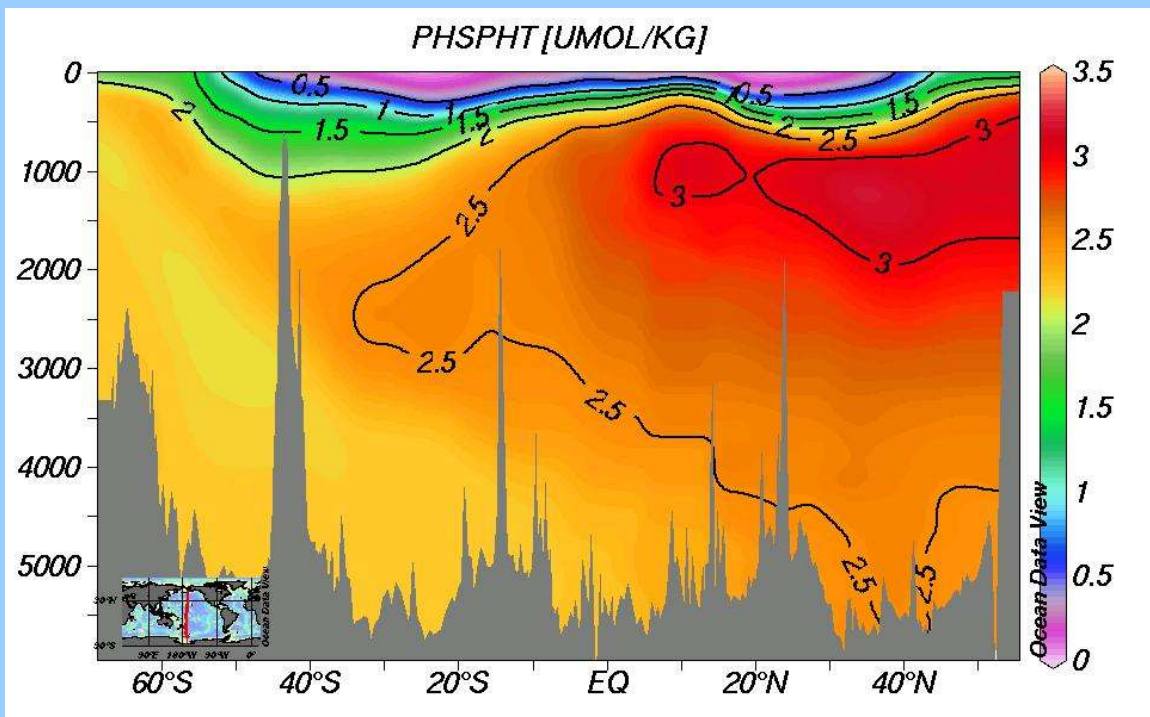
- **soft tissue pump**





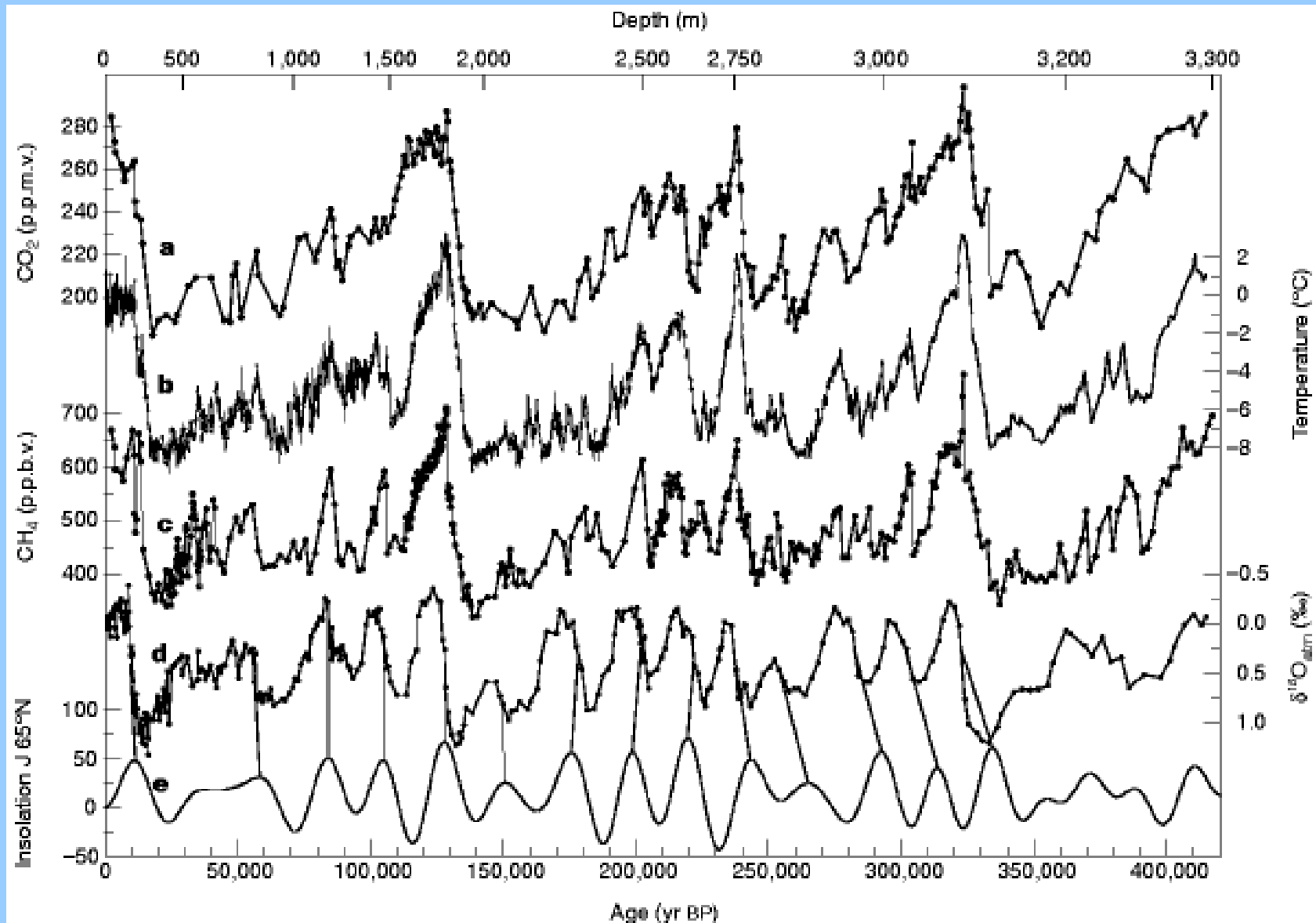
## DIC and PO<sub>4</sub>

- common biological processes
- **soft tissue pump**





# Could change in ocean carbon cycle cause glacial-interglacial changes in atmospheric pCO<sub>2</sub>?



# Could changes in the oceans have reduced atmospheric $p\text{CO}_2$ by $>80\text{ppmv}$ at the Last Glacial Maximum?

## *Changes in solubility pump:*

Cool deep ocean by $2^\circ\text{C}$	-20ppmv
Cool upper ocean (thermocline) by $5^\circ\text{C}$	-20ppmv

## *Changes in biological pumps:*

Extreme enhancement of biological export efficiency	-40ppmv
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## *Changes to circulation and mixing?*

STILL AN OPEN QUESTION



